#### STATE OF ARKANSAS

# Arkansas Geological and Conservation Commission Norman F. Williams, Geologist-Director

## WATER RESOURCES CIRCULAR No. 2

# GROUND-WATER RESOURCES IN A PART OF SOUTHWESTERN ARKANSAS

Вy

H. B. Counts, D. B. Tait, Howard Klein, and G. A. Billingsley

U. S. GEOLOGICAL SURVEY



Prepared in Cooperation with the U. S. Geological Survey and the University of Arkansas Institute of Science and Technology

Little Rock
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## GROUND-WATER RESOURCES IN A PART OF SOUTHWESTERN ARKANSAS

H. B. Counts, D. B. Tait, Howard Klein, and G. A. Billingsley

#### ABSTRACT

This report gives information about the geography, geology, and ground-water resources of an area about 3,900 square miles located in southwestern Arkansas. The topography consists of low rounded hills or low ridges trending in a general easterly direction with some terraces and flat bottomland along the major streams. The area is drained by the Red River, Little River, and the Little Missouri River. The climate is mild, with an average growing season of 226 days and the average annual precipitation ranges from about 45 to 50 inches. Agriculture is the principal industry.

Formations exposed in the area range from Paleozoic to Quaternary in age. Formations of Cretaceous age have a general southerly or southeasterly dip and are exposed as eastward or northeastward trending bands through the report area. Sediments of Tertiary age overlie the Cretaceous deposits in the southern part and rocks of Paleozoic age crop out in the northern part. Deposits of Quaternary age overlie the older formations, either as terraces or alluvial deposits in the major stream valleys.

Ground water is used in the area at an average rate of about 8.5 million gallons per day of which approximately 3 million gallons per day are used by municipalities. Ground water is wasted at an estimated rate of about 2 million gallons per day. The total combined use for domestic, industry, and irrigation is about 3.5 million gallons per day.

There is a considerable difference in the quality of water from the different waterbearing formations as indicated by the analyses of samples from 328 wells. Also, the water in the Cretaceous formations is progressively more mineralized down dip.

The most important water-bearing units are the Trinity group, the Tokio formation, and the Nacatoch sand, all of which are of Cretaceous age.

## GROUND-WATER RESOURCES IN A PART OF SOUTHWESTERN ARKANSAS

#### INTRODUCTION

This is one of a series of reports on the ground-water resources of Arkansas made under a program of investigations with the United States Geological Survey, cooperating with the University of Arkansas, Bureau of Research and Institute of Science and Technology from 1946 to 1950, the Arkansas Resources and Development Commission from 1950 to present, and with the University of Arkansas, Agricultural Experiment Station.

The areas in Arkansas covered by these reports are shown on Figure 1. Other reports giving information about ground water in Arkansas are Hale and others (1947), Branner (1937), Stephenson and Crider (1916), and Veatch (1906). The last report is particularly complete for southwest Arkansas to the date of its publication. These and other references will be found at the end of this report.

The investigation upon which this report is based was begun in the summer of 1950. The field work was started by Howard Klein in August 1950 and was completed by H. B. Counts and D. B. Tait from March to September of 1951. The work was under the direct supervision of R. C. Baker, District Geologist, and under the general direction of A. N. Sayre of the U. S. Geological Survey. Chemical analyses of water samples were made by the U. S. Geological Survey.

#### PURPOSE AND SCOPE OF INVESTIGATION

The purpose of this report is to present information about the occurrence and quality of ground water in part of southwestern Arkansas, with particular reference to the deposits of Cretaceous age. Records of 397 wells and chemical analyses of water from 328 wells are listed. Maps are given showing well locations, the outcrop areas of and quality of water in the Tokio formation and the Nacatoch sand, the outcrop area for the Trinity group and location of wells tapping the Trinity, Woodbine formation, Brownstown marl, Ozan formation, Wilcox formation, and deposits of Quaternary age.

#### LOCATION OF AREA

The area covered in this report consists of about 3,900 square miles located in the southwestern part of Arkansas. It includes all of Little River, Hempstead and Sevier Counties, the southern part of Howard, Pike and Clark Counties, and the northern part of Miller and Nevada Counties. The area is roughly rectangular in shape, with Texarkana near the southwest corner, Rosston near the southeast corner, Arkadelphia near the northwest corner, and DeQueen near the northwest corner.

#### **ACKNOWLEDGMENTS**

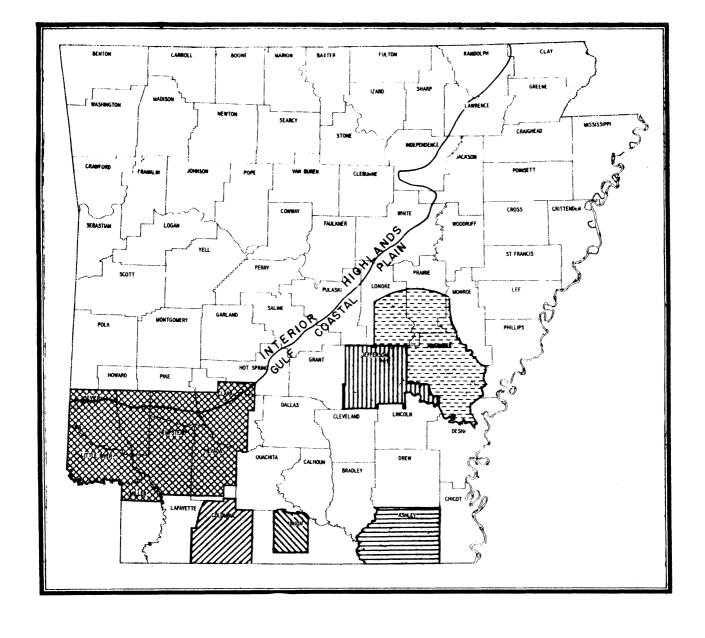
The writers are grateful to all persons who gave information or help in the investigation, including Clyde Zinn and Dennis Bell of Hope Light and Water Co.; F. E. Wells, Wallace Elliot, Howard Hipp, R. H. Herron, Milton Garrison, and Erman Williamson, water well drillers; K. Scott of Carter Oil Co.; L. Bartells of McAlester Fuel Oil Co.; Nonus Harris and H. Harris, guides.

#### WELL-NUMBERING SYSTEM

The well-numbering system used in this report is based upon the location of the wells with respect to the federal land-survey used in Arkansas. The component parts of a well number are the township number, the range number, the section number, and three lower-case letters which indicate respectively, the quarter section, the quarter-quarter section, and the quarter-quarter-quarter section in which the well is located. The lower-case letters are assigned in counter-clockwise order beginning with a in the northeast quarter or quarter-quarter or quarter-quarter section. Serial numbers are appended to each well located within the quarter-quarter-quarter section.

This system of numbering wells according to their location is illustrated in Figure 2.

Map numbers were also assigned to each well as they were tabulated in Table 11. It was thought that the regular well numbers would be hard to read on the maps, especially where there



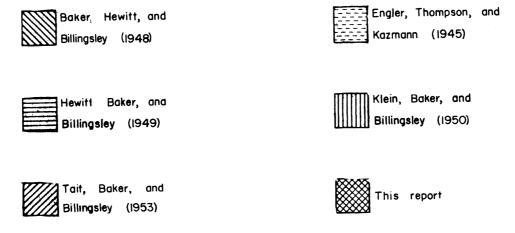


Figure 1. Areas in Arkansas Covered by Recent Ground-Water Reports.

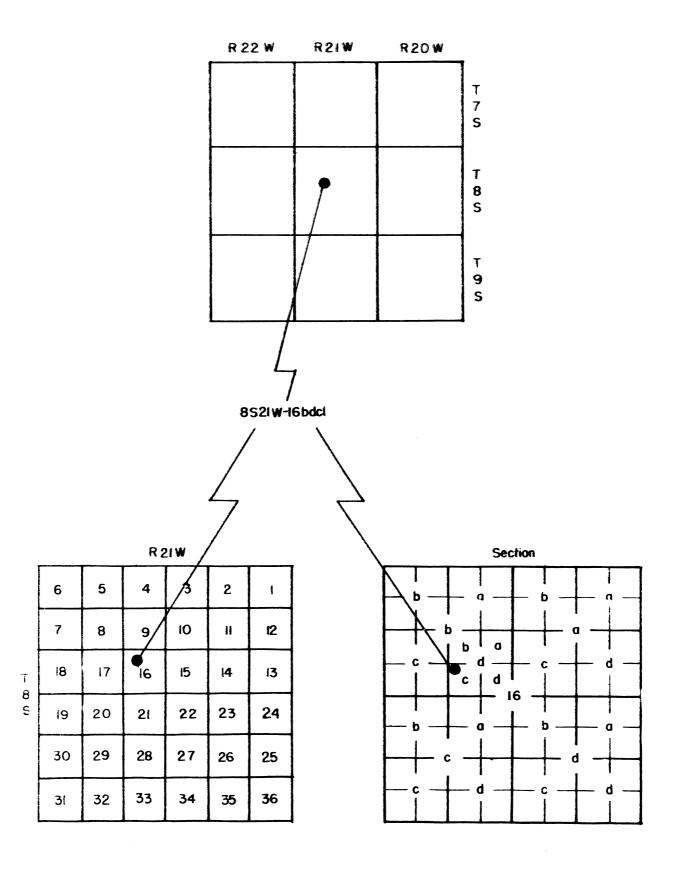


Figure 2. Sketch Showing Well-Numbering System.

are a large number of wells in a small area, and that the shorter map number would be easily referred to in the text of this report.

#### **GEOGRAPHY**

#### Topography and Drainage

The area lies entirely in the Gulf Coastal Plain except for a small part on the north which lies in the Interior Highlands.

The land surface has a general slope to the south, the highest elevations being about 800 feet above sea level for hill tops at the northern edge of the area and the lowest elevations being about 165 feet in the flood plains of streams near the southern boundary of the area. Elevations in the inter-stream areas average from 300 to 600 feet above sea level.

About 2,500 of the 3,900 square miles of the area covered by this report lie directly on the outcrop area of the formations of Cretaceous age. The surface expression of the Cretaceous area is characterized by low rounded hills, aligned in a general northeastward direction, the trend conforming to the outcrop bands of the formations. The more resistant formations, such as the Nacatoch sand and the Tokio formation, stand up as low ridges.

The larger streams, such as the Little, Saline, Little Missouri. Ouachita. and Red rivers flow to the south or southeast across the area from the upland of Paleozoic rocks. They have breached the series of alined hills and ridges and have eroded their valleys to about 100 or 200 feet below the higher hills. The major streams have broad, flat alluvial valleys which range from one to six miles wide, generally being widest at the confluence of major streams or where the streams cross less resistant formations, at which places alluviation is caused by a decrease in velocity. The largest alluviated area is in the southwest part of the area in the valley of the Red and Little rivers. Where the streams have become sluggish and meander over a wide, flat area, frequent floods occur. The bottomlands and backwater areas may remain swampy for extended periods of time after flooding.

The alluvial valleys of the streams are at many places flanked by older terrace deposits. The altitude of the highest noted terrace in the area is about 500 feet above sea level and it occurs about 10 miles east of Nashville in an old valley of the Little Missouri River. The surface slopes southward and drops to an altitude of 470 feet in a distance of 4 miles. Occasional patches of gravel may indicate older terraces. Near Prescott vounger terraces occur about 100 feet above the present flood plain of the Little Missouri River. A wide terrace north and south of Ashdown in the southwestern part of the area lies at about the same elevation and was probably formed at the same time. A terrace along the Red River east and south of Fulton stands at about 60 feet above the bottomland of the river.

#### Climate

Growing season and rainfall information for selected weather stations in the area are shown on Tables 1 and 2. Growing season, temperature and rainfall information for Hope, near the center of the area, are shown on Figure 3. The average frost-free growing season varies over the area with a maximum of about 233 days at Hope and a minimum of 216 days reported for Index, in Little River County, with an overall average of 226 days. The average annual precipitation varies from 44 inches to 51 inches for an 80 year period. The maximum and minimum over the area are about 60 to 40 inches respectively.

In general, the summers are hot and the winters are mild. Rainfall is plentiful for most agri-

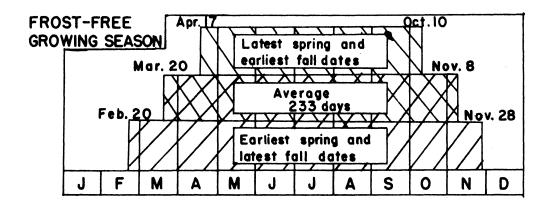
TABLE 1

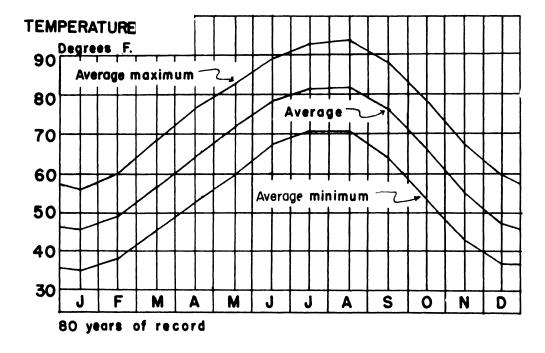
Average Dates of Earliest and Latest Killing Frosts and
Lengths of Frost-Free Growing Seasons for
Selected Weather Stations

	Average date of last killing frost	Average length of growing season (days)	Average date of first killing frost
Норе	March 20	233	Nov. 8
Index	March 28	216	Oct. 30
Prescott	March 24	228	Nov. 7
Texarkana	March 21	232	Nov. 8
White Cliffs	March 26	225	Nov. 6
Average	March 24	226	Nov. 5
Texarkana White Cliffs	March 21 March 26	232 225	Nov. 8 Nov. 6

TABLE 2
Average Monthly Precipitation for Selected Weather Stations

Station	Jan.	Feb.	Mar.	April	May	June	July	Aug.	Sept.	Oct.	Nov.	Dec.	Annual Average
Arkadelphia	4.63	3.82	4.88	5.23	5.20	4.49	3.71	3.65	2.77	3.09	3.85	5.04	50.36
Fulton	4.57	3.28	4.55	4.60	4.51	3.81	3.56	2.82	2.49	2.58	3.79	4.10	44.66
Hope	4.99	4.07	5.08	5.31	5.09	3.74	4.35	3.28	3.06	3.19	4.65	4.64	51.35
Index	4.28	3.51	4.56	4.65	4.31	3.21	3.91	2.97	2.47	2.86	4.24	4.66	45.63
Prescott	4.74	3.80	4.56	5.12	4.89	3.72	3.54	3.16	2.92	2.63	3.71	4.66	47.45
Texarkana	4.48	3.37	4.30	4.94	4.52	3.71	4.11	3.37	2.67	3.41	3.67	4.61	47.16
White Cliffs	3.96	3.44	4.47	5.47	4.81	3.51	2.69	2.90	2.51	3.13	3.33	4.64	44.86





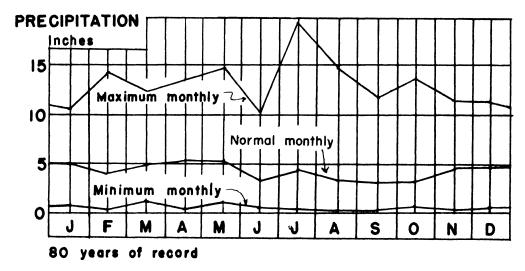


Figure 3. Selected Climatological Data for Hope, Ark.

cultural needs with the largest amounts falling in the winter and spring months. Extended droughts are unusual but occur occasionally. The summer of 1951 brought a fairly long drought which was not too damaging to crops, but the summer of 1952 brought a damaging drought.

#### Cultural Development

The population of the area has been variable. In general it increased up to about 1920, was about constant from 1920 to 1940, and has decreased since 1940. The decrease in population during the last few years can be attributed to the general migration from farms to industry during the war years, to a decrease in cotton acreage, and to an increase in the beef cattle industry which takes less labor than other types of farming and also tends to combine several small farms into larger ones.

Most of the people live on farms and are employed in agriculture. The towns generally have grown in size and some of them have had remarkable increases in population.

The population of counties and of principal municipalities is given in Tables 3 and 4 respectively.

Farming is the most important occupation of the area. The soils of the area range from heavy clay to light sandy loams. A large variety of crops are grown in the area and several places are noted for specific crops, such as Hope for watermelons, and Nashville for peaches. The beef cattle industry and the production of corn and hay have increased during the past few years. Cotton production generally has decreased. Radishes, strawberries, and cucumbers are other important crops.

Lumbering, although not a large industry, is a stable one and is of long standing.

The only cement plant in Arkansas is located about 12 miles south of Nashville where there are large reserves of raw material.

#### GENERAL GEOLOGY

Published reports by Miser and Purdue (1929) and Dane (1929) give information about the geology of the area, particularly with reference to the surface geology. The following statements about the geology are taken largely from those reports.

Rocks of Paleozoic age crop out along the northern part of the area. The formations at

TABLE 3
Population of Counties by Decades from 1890 to 1950¹

County	1950	1940	1930	1920	1910	1900	1890
Clark	22,958	24,402	24,932	25,632	23,686	21,289	20,997
Hempstead	25,045	32,770	30,847	31,602	28,252	24,101	22,796
Howard	13,360	16,621	17,489	18,565	16,898	14,076	13,789
Little River	11,232	15,932	15,515	16,301	13,597	13,731	8,903
Miller	31,382	31,874	30,586	24,021	19,555	17,558	14,714
Nevada		19,869	20,407	21,934	19,344	16,609	14,832
Pike	9,995	11,786	11,792	12,397	12,565	10,301	8,537
Sevier	12,244	15,248	16,364	18,301	16,616	16,339	10,072
Totals	140,977	168,502	167,932	168,753	150,513	134,004	114,640

<sup>&</sup>lt;sup>1</sup>The total county population is given in this table. Only three counties, Hempstead, Little River and Sevier, are completely within the area covered by this report.

TABLE 4
Populations of Principal Municipalities by Decades from 1890 to 1950

Municipality	1950	1940	1980	1920	1910	1900	1890
Arkadelphia <sup>1</sup>	6,819	5,078	3,380	3,311	2,745	2,739	2,455
Ashdown <sup>a</sup>	2,738	2,332	1,607	2,052	1,247	400	
DeQueen*	3,015	3,055	2,938	2,517	2,018	1,200	
Dierks1	1,253	1,459	1,544	1,495	272		
Foreman <sup>a</sup>	907	1,007	1,056	1,408	612		
Gurdon'	2,390	2,045	2,172	1,469	1,284	1,045	802
Hope <sup>2</sup>		7,475	6,008	4,790	3,639	1,644	1,937
Murfreesboro*		835	733	730	516	200	159
Nashville <sup>1</sup>	3,548	2,782	2,469	2,144	2,374	928	810
Prescott <sup>2</sup>	3,960	3,177	3,033	2,691	2,705	2,005	1,287
Texarkana*	15,875	11,821	10,764	8,257	5,655	4,914	3,528

<sup>&</sup>lt;sup>1</sup> Uses surface water.

<sup>&</sup>lt;sup>2</sup> Uses ground water.

<sup>\*</sup> Uses surface water and ground water.

the surface consist of the Stanley shale, the Jackfork sandstone, and the Atoka formation. The rocks extend to the south but they are overlain by younger formations. The contact with the younger formations dips in a general south or southeasterly direction at a rate of about 120 feet per mile.

Over most of the area the Paleozoic rocks are overlain by sediments of Cretaceous age. In the southern part of the area rocks of Jurassic age occur between the Paleozoic rocks and the Cretaceous formations but they are deeply buried and do not crop out in Arkansas.

The deposits of Cretaceous age have the form of a wedge, thinning to a feather edge to the north against the Paleozoic rocks and thickening rapidly to the south. Stratigraphically the Cretaceous deposits have been divided into the Lower Cretaceous (Comanche) series and the Upper Cretaceous (Gulf) series. A generalized geologic column of the area covered by this report is given in Table 5. Plate 1 shows the geologic sequence and the dip of the Cretaceous formations and Plate 2 shows the surface geology of the area.

The Lower Cretaceous deposits crop out in a wedge shaped east-west trending band bordering the Interior Highlands. About 10 miles east of Murfreesboro the outcrop is about 3 miles wide and reaches a width of more than 12 miles about 4 miles east of DeQueen. The dip of the strata averages about 80 to 120 feet per mile southward.

The majority of the Lower Cretaceous deposits in Arkansas are near-shore gravel, sand, clay, and lime, parts of which have become indurated. To the south there is a change in facies from a shallow water or marginal environment to a deeper water type material. The thickness of the Lower Cretaceous strata ranges from about 500 feet at the outcrop to more than 6,000 feet in northern Louisiana.

The only major outcrop area of the Lower Cretaceous strata in Arkansas is of beds of the Trinity group. Of minor importance are small outcrop areas of the Goodland limestone and Kiamichi formation, located at the Arkansas-Oklahoma boundary just south of the place where the Little River enters Arkansas from Oklahoma. These formations are younger than the Trinity.

Deposits of Upper Cretaceous series overlie Lower Cretaceous deposits. The position of Upper Cretaceous strata in relation to the Lower Cretaceous suggest a shift in the direction of the shore line during the times of flooding, for instead of striking parallel to older strata the formations seem to pivot to northeast-southwest. Such a discordance in strike could be the result of eastward tilting and downwarping of the Paleozoic floor (Veatch 1906). The downwarping was probably continuous during late Cretaceous time since the older formations at their outcrops became thinner to the northeast. However, it is recognized that an eastward thinning of these deposits could have been the result of more rapid erosion or of lesser deposition in this area than in the west. The beds dip to the south and southeast at rates of from about 40 to 80 feet per mile.

Igneous rocks represented by peridotite plugs are exposed in 4 places about 3 miles southeast of Murfreesboro in Pike County. The peridotite is generally believed to be similar in character and mode of occurrence to that of South Africa. Many diamonds have been taken from 3 of these intrusions. The peridotite is considered to be of early late Cretaceous age (Miser and Purdue 1929).

The deposits of the Cretaceous system have been divided into 12 groups or formations and they are described in this report in connection with their water-bearing properties.

Along the southern edge of the area sediments of Tertiary age overlie the Cretaceous deposits. These deposits yield water to wells mainly in Little River, Miller and Nevada Counties. They consist of the Midway formation of the Paleocene series and the Wilcox formation of the Eocene series. The Wilcox yields water to wells in the southern part of the area and is described in connection with its water-bearing properties.

Deposits of Quaternary age overlie the older formations occurring as terraces and in the bottomlands along the major stream valleys. These deposits yield water to wells mainly in Little River, Miller and Nevada Counties. They are described later in connection with their water-bearing properties.

#### GROUND WATER

#### Occurrence and Movement

Only a brief discussion of the principles of occurrence and movement of ground water will be given in this report. A comprehensive description of these principles is given by Meinzer (1923) and most of the following discussion is of their application to the Southwestern Arkansas area.

TABLE 5
Generalized Geologic Column of Deposits in Southwestern Arkansas

Era	System	Series	Group or formation	Thickness (feet)	Character of materials	Water supply				
o i c	Quaternary.		Alluvium and terrace deposits.	0-90	Sand, gravel, silt, and clay.	Yields moderate supplies of hard water to domestic wells and to public supply wells at Ashdown and Foreman.				
20		Eocene.	Wilcox formation.	0-600	Lignitic sand and clay.	Yields water to domestic wells and to one public supply well at Rosston.				
Cen	Tertiary.	Paleocene.	Midway formation.	0-600	Clay, blue plastic. Clay, blue fossiliferous, bottom 30 to 50 feet.	Does not yield water to wells.				
				0-150	Clay, fossiliferous, calcareous with interbedded limestone. Shale, blue in subsurface.	Does not yield water to wells.				
	G u 1 f)	Nacatoch sand.	0-500	Sand massive cross-bedded, limestone lenses and calcareous clay.	Yields moderate supplies of good quality water to domestic, industrial, and public supply wells on outcrop area and a distance of 2 to 15 miles downdip to the south.					
		<b>s</b>	Saratoga chalk.  -Unconformity.	0-60	Chalk, hard white with inter- bedded blue marl. Fossiliferous.	Does not yield water to wells.				
		n o	Marlbrook marl.	0-200	Marl, blue to gray, fossiliferous.	Does not yield water to wells.				
		O U	Annona chalk.	0-100	Chalk, massive white.	Does not yield water to wells.				
		Creta	-Unconformity.  Ozan formation.	0-250	Clay, bluish to tan and clay and marl, sandy. Sand glauconitic at base from 0-20 feet thick.	Yields small amounts of very highly mineralized water to wells.				
	•	H O	Brownstown marl.	0-200	Clay, gray to tan, fossiliferous calcareous.	Yields small amounts of highly mineralized water to wells.				
s o i c	o da	<b>d</b> d	Tokio formation.	0-350	Sand, cross-bedded quartz, gravels and clay, gray lignitic.	Yields moderate supplies of good quality water to domestic, industrial, and public supply wells. Locally water has high iron content.				
0	et et	omanche)	-Unconformity.  Woodbine formation.  -Unconformity.	0-250	Clay, red and gray, sand and gravel, yellowish cross-bedded.	Yields small amounts of poor quality water to a few domestic wells.				
×	0 0		_ <b>_</b> _	c h	c h	c p	Kiamichi formation.	0-20	Clay, soft gray, marl fossiliferous and lenses of limestone. Present in very small area in Little River County.	Does not yield water to wells.
			Goodland limestone.	0-50	Limestone gray sandy and clay gray. Present in very small area in Little River County.	Does not yield water to wells.				
		) 8		0-900	Sand, fine, white, interbedded with red and clay, and limestone.	Yields moderate supplies of fair quality water in Howard and Sevier counties.				
		n o e o		0-100	Limestone, interbedded gray, clay and gypsum, gray.	Does not yield water to wells.				
	Lower Cretac	Trinity group.	0-40	Gravel.	Yields moderate supplies of fair quality water in DeQueen and vicinity.					
		D H	group.	0-400	Clay, red, interbedded with gray sand.	Does not yield water to wells.				
		•		0-40	Limestone interbedded with dark shale.	Does not yield water to wells.				
		- 1		0-50	Gravel.	Yields moderate supplies of fair quality water in DeQueen and vicinity.				
	Jurassic.			0-?	Red beds and anhydrite.	Not a source of fresh water.				
Paleozoic	Mississippian and Pennsylvanian		Atoka formation. Jackfork sandstone. Stanley shale.	0-7	Sandstone and shale, highly folded.	Yields small amounts of hard water to domestic wells.				

Ground water is derived chiefly from water that falls as rain or snow. A part of the precipitation runs off in streams, a part is returned to the atmosphere by evaporation and/or by transpiration from vegetation, and a part sinks into the zone of saturation in which all of the openings of the rocks are filled with water.

In most places ground water is slowly but steadily moving under the influence of gravity from areas of intake to areas of discharge. In the more permeable rocks, such as coarse sand, gravel and porous limestone, the water moves with comparative freedom although the movement is very slow, as compared to the flow of a stream. Such rocks are capable of yielding abundant supplies of water to wells. In less permeable rocks, such as shale or clay, molecular attraction and surface tension retard the movement of the water and the movement may be almost infinitely slow. Such rocks yield little or no water to wells.

On the outcrop of water-bearing beds, the water is usually unconfined and does not rise in wells above the water table, which is the upper surface of the zone of saturation and the level at which water is first encountered.

Where water-bearing beds are inclined and are enclosed between relatively impermeable strata, the water is under artesian pressure and will rise in wells above the level at which it is struck. If the altitude to which the water will rise is greater than the altitude of the land surface, flowing wells will result.

The deposits of Cretaceous age consist chiefly of clay, shale, limestone, chalk and marl interbedded with sand and at some places gravel. They dip south and southeastward. The land surface also slopes in the same direction. Hence, artesian conditions exist in all parts of the area, and in many places, mainly in the river valleys, flowing wells are obtained.

Water gets into the formations of southwestern Arkansas by percolation of surface water through the permeable portions of the outcrop areas. The locations of the outcrop areas for the main aquifers in southwestern Arkansas are shown on Plates 3, 4, and 5. Water moves from these areas down the dip of the formations which is south, southeast. On the basis of a few pumping tests and an approximation of the gradient, the rate of ground-water movement is on the order of about 5 feet per year. This figure is not exact but it does illustrate the slow rate at which the water is moving.

The water levels in most wells are subject to fluctuations of varying magnitude. These fluc-

tuations are due to many different causes, but the larger ones are generally a result of a change in the ratio between the rate of ground-water intake or recharge and the rate of loss or discharge. Most water-table wells are supplied in part from intake areas close at hand and respond to changes in rainfall with only a moderate lag. In shallow wells the water level may rise several feet after heavy rains and then decline during dry periods. During prolonged droughts such wells may go dry. The water levels in artesian wells, located at considerable distances from the outcrop area of the water-bearing beds, are not affected immediately or strongly by seasonal or yearly changes in rainfall. Instead there is a time lag and the magnitude of the fluctuations is not as great. Large fluctuations may be caused by a series of wet or dry years. Large fluctuations in pressure in such wells and the accompanying rise and fall in water levels are generally caused by withdrawals of groundwater from the wells themselves or from other wells in the vicinity. Minor fluctuations are caused by changes in barometric pressure, earth tides, passing trains, and other temporary loads upon the aquifer.

When a well is pumped or allowed to flow the water level in the well declines, a hydraulic gradient is developed toward the well from all directions, and water flows toward the well. Within limits the amount of water that will enter a well varies directly with the amount the water-level is lowered. The difference between the static (pre-pumping) water level and the pumping water level in the well is called the drawdown.

Large withdrawals of ground water are almost certainly to be accompanied by a general lowering of the water table or artesian pressure. A cone of depression gradually spreads out in all directions from the center of pumping until large areas may be affected. However, this is usually not very serious unless the rate of decline persists or the trend is such as to indicate that the pumping lift may eventually exceed the economic limit, or that the water-bearing beds will be unwatered.

## PERMEABILITY OF WATER-BEARING MATERIALS

The rate of movement of ground water is determined by the size, shape, number, and degree of interconnection of the interstices, by the density and viscosity of the water, and by the hydraulic gradient. The capacity of a water-bearing material for transmitting water under a hydraulic gradient is known as its permeability. Meinzer's coefficient of permeability may

be expressed in field terms as the number of gallons of water a day, at 60°F., that is conducted laterally through each mile of the waterbearing bed under investigation (measured at right angles to the direction of flow), for each foot of thickness of the bed and for each foot per mile of hydraulic gradient (Stearns 1928).

The field coefficient of permeability, as determined from aquifer tests, is the same as Meinzer's coefficient defined above except that it is at the existing temperature. This is the coefficient of permeability that is generally used by the U. S. Geological Survey.

The coefficient of transmissibility is a similar measure for the entire thickness of the waterbearing formation and may be defined as the number of gallons of water a day, under prevailing temperature, transmitted through each 1-mile strip extending the height of the aquifer under a hydraulic gradient of 1 foot per mile. It is the field coefficient of permeability multiplied by the thickness of the aquifer in feet.

The coefficient of permeability of the Nacatoch sand and of the Tokio formation at Hope, in Hempstead County, were determined by aquifer tests on wells 309 and 207 of the Hope Water and Light Co. Measurements of discharge were made with permanent meters in the discharge lines and measurements of drawdown and recovery were made with a steel tape. The recovery formula of Theis (1935), which was used in computing transmissibility at each well, may be expressed as follows:

$$T = \frac{264 \ Q}{s} \log_{10} \frac{t}{t^1}$$

T = coefficient of transmissibility, in gallons per day per foot

Q = pumping rate, in gallons per minute

t = time since pumping began, in minutes

 $t^1$  = time since pumping stopped, in minutes

s = residual drawdown at the pumped well, in feet at time t1.

The residual drawdown (s) is computed by subtracting the static water-level measurement (Tables 6 and 7) from water-level measurements made after pumping stops (Figure 8).

$$\log \frac{t}{t^1}$$

 $\frac{\log \frac{t}{-t^1}}{t^1}$  The proper ratio  $\frac{t}{s}$  is determined graph-

ically by plotting  $log_{10} \frac{t}{t^1}$  on the logarithmic coordinate of semi-logarithmic paper as shown on

TABLE 6 Data on Aquifer Test on Well 309 in the Nacatoch Sand

Figures 4 and 5.

Time since pumping started (minutes)	Time since pumping stopped (minutes) t <sup>1</sup>	$\frac{t}{t^1}$	Depth to water level (feet)	Residual drawdown (feet) (s)	Remarks
	-	The second second second second second	205.22	A	Static water level
0					Pump started
<b>22</b> 0	0				Pump stopped
228	8	28.5	234.48	29.26	
230	10	<b>2</b> 3	232.39	27.17	
233	13	17.9	230.24	25.02	
236	16	14.7	228.54	23.32	
239	19	12.6	226.98	21.76	
242	22	11.0	225.70	20.48	
<b>25</b> 0	30	8.33	223.37	18.15	
255	35	7.3	221.92	16.70	
260	40	6.5	220.83	15.61	
267	47	5.69	219.69	14.47	
272	<b>52</b>	<b>5.2</b> 3	<b>218.99</b>	13.77	
280	60	4.67	217.90	12.78	
300	80	3.75	215.92	10.70	
310	90	3.45	215.17	9.95	
325	105	3.10	214.19	8.97	
340	120	2.83	213.45	8.23	
359	139	2.57	212.64	7.42	
370	150	2.46	212.17	6.95	
385	165	2.33	211.68	6.46	
388	168	2.30	211.60	6.38	
406	186	2.18	<b>211.0</b> 9	5.87	
418	198	2.10	210.82	5.60	
433	213	2.03	210.51	5.29	
445	225	1.98	210.28	5.06	
460	240	1.92	210.04	4.82	

TABLE 7
Data on Aquifer Test on Well 207 in the Tokio Formation

Time since pumping started (minutes) t	Time since pumping stopped (minutes) t <sup>1</sup>	t t1	Depth to water level (feet)	Residual drawdown (feet) (s)	Remarks
-		de analis de la cista et Médide agridige a	66.92		Static water level
0	_				Pump started
440	0				Pump stopped
441	1	441	88.95	22.03	
442.5	2.5	177	84.20	17.28	
443.5	3.5	127	82.52	15.60	
445.5	5.5	81	80.63	13.71	
447	7	64	78.77	12.85	
450	10	45	78.08	11.16	
452	12	37.6	77.33	10.41	
455	15	30	76.49	9.57	
457	17	27	76.03	9.11	
461	21	22	75.27	8.35	
465	25	18.5	74.66	7.74	
472	32	14.7	73.88	6.96	
482	42	11.5	73.05	6.13	
492	52	9.5	72.46	5.54	
503	63	8	71.94	5.02	
516	76	6.8	71.47	4.55	
530	90	5.9	71.06	4.14	
550	110	5	70.60	3.68	
572	132	4.3	70.21	3.29	
615	175	3.5	69.65	2.73	
705	265	2.66	68.91	1.99	
770	330	2.34	68.57	1.65	
800	360	2.22	68.43	1.51	
915	475	1.92	68.07	1.15	

The average discharge (Q) of well 309 in the Nacatoch sand was 280 gallons per minute.

When values for s,  $\log_{10} \frac{t}{t_1}$ , and Q are substituted

in the Theis recovery formula (Figure 4), the coefficient of transmissibility of the Nacatoch sand near the pumped well is found to be about 3,600 gallons per day per foot. The coefficient of permeability, which is determined by dividing the coefficient of transmissibility by the thickness of the water-bearing material, m, (230 feet), was found to be about 16 gallons per day per square foot. This figure may be too low because well 309 has 230 feet of open hole in the Nacatoch sand which contains some shale and clay that are not water-bearing. The effective thickness in this well is probably about 100 feet which would give a coefficient of permeability of 36.2 gallons per day per square foot.

The average discharge of well 207 in the Tokio formation was 184 gallons per minute. When

values for s,  $\log_{10} \frac{t}{t^1}$ , and Q are substituted in

the Theis recovery formula (Figure 5) the coefficient of transmissibility of the sand layer in the Tokio formation, near the pumped well, is found to be about 4,500 gallons per day per foot.

This figure may seem high when compared to 1,300 gallons per day per foot obtained from a test on well 175 in the same formation as shown on the following page. Although no samples are available to make comparisons, it is believed that the difference reflects the true character of the materials because drillers report that the Tokio formation is much coarser in the vicinity of Hope than in Howard County where well 175 is located.

The late points on Figure 5 indicate that there is probably leakage from some sand above or below into the aquifer, therefore, these were not used when drawing the line for calculations of the transmissibility.

The thickness of the sand is 49 feet (see drillers log 207), and, therefore, the coefficient of permeability is about 90 gallons per day per square foot.

Another aquifer test in the Tokio formation was made at the Ideal Cement Company, in Howard County. Well 175 was pumped at 70 gallons per minute and water-level measurements were made in well 177. Table 8 shows the drawdown in well 177 with well 175 pumping. The coefficient of transmissibility and the coefficient of storage were determined by the Theis nonequilibrium method. The coefficient of storage is defined as the amount of water, in

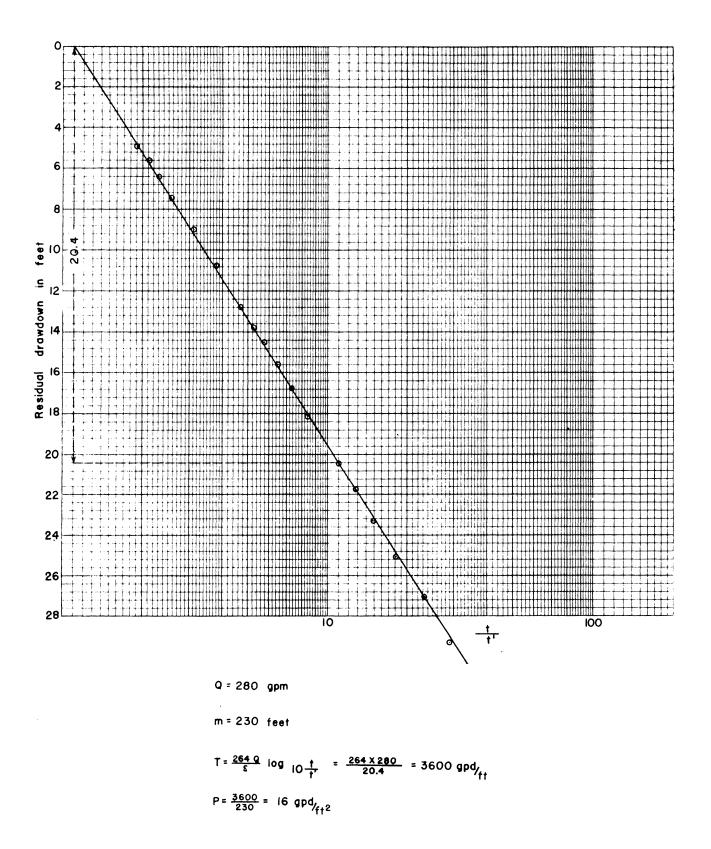
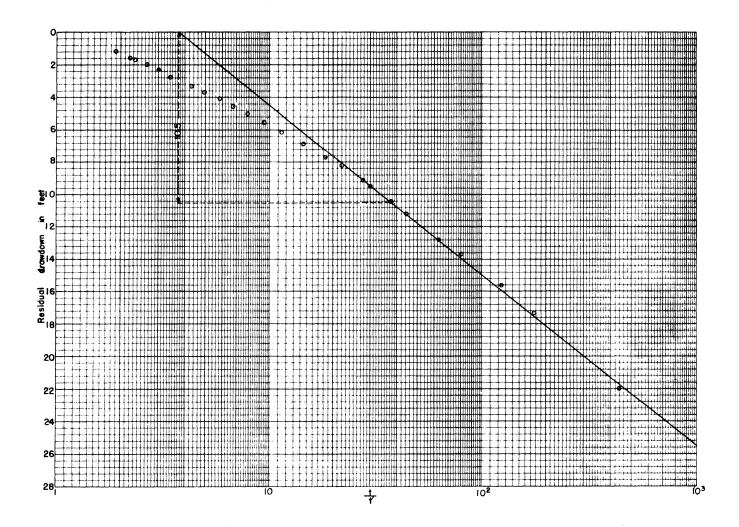


Figure 4. Curve for Aquifer Test on Well 309 in the Nacatoch Sand.



$$Q = 184 \text{ gpm}$$

$$P = \frac{4500}{49} = 90 \text{ gpd}_{41}^2$$

Figure 5. Curve for Aquifer Test on Well 207 in the Tokio Formation.

TABLE 8

Data on Aquifer Test in the Tokio Formation with Well 175 Pumping 70 gpm and with

Water-Level Measurements Made in Well 177

Time since pumping started (minutes)	r 2,080 feet distance from pumping well r <sup>2</sup>	Depth to water level (feet)	Drawdown (feet)	Remarks
t	t (min)			
		115.95		Static water level
0				Pump started $Q=70$ g.p.m.
22	$1.97 \times 10^{5}$	115.95	0	• • • • • • •
53	8.15 x 10 <sup>4</sup>	115.96	0.01	
69	6.27 x 10 <sup>4</sup>	115.96	0.01	
84	5.15 x 10 <sup>4</sup>	115.98	0.03	
98	4.41 x 10 <sup>4</sup>	115.99	0.04	
123	3.5 x 10 <sup>4</sup>	116.03	0.08	
176	2.46 x 10 <sup>4</sup>	116.17	0.22	
193	2.24 x 10 <sup>4</sup>	116.22	0.27	
207	2.09 x 10'	116.27	0.32	
230	1.88 x 10 <sup>4</sup>	116.36	0.41	
275	1.57 x 10°	116.59	0.64	
308	1.4 x 10'	116.76	0.81	
355	1.22 x 10 <sup>4</sup>	117.03	1.08	
405	1.07 x 10'	117.30	1.35	
455	$9.5 \times 10^{\circ}$	117.60	1.65	
505	$8.55 \times 10^{3}$	117.87	1.92	
555	$7.8 \times 10^{3}$	118.15	2.20	
620	$6.96 \times 10^{3}$	118.48	2.53	
675	$6.4 \times 10^{3}$	118.87	2.92	
755	$5.72 \times 10^{3}$	119.27	3.22	
840	$5.15 \times 10^{3}$	119.59	3.64	
925	$4.67 \times 10^{3}$	119.95	4.00	
1015	$4.25 \times 10^{3}$	120.33	4.38	
1110	$3.89 \times 10^{3}$	120.70	4.75	
1205	$3.58 \times 10^{3}$	120.91	4.96	
1325	$3.26 \times 10^{3}$	121.29	5.34	Pump stopped

cubic feet, that will be discharged from a vertical column of the aquifer having a basal area of one square foot when the water level is lowered one foot (Theis 1935). The Theis nonequilibrium formula may be written as follows:

$$s = \frac{114.6q}{T} \int_{\frac{1.87r^3S}{T_t}}^{\infty} \frac{e^{-u}}{u} du$$
 (1)

Where:

s - drawdown of water level, in feet.

Q = discharge of pumped well, in gallons per minute.

r = distance of observation well from pumped well, in feet.

T = coefficient of transmissibility, in gallons per day per foot.

S = coefficient of storage.

t = time well has been pumped, in days.

The exponential integral of formula (1) is replaced by the term W(u) which is read "Well function of u", and the equation is rewritten as follows:

$$s = \frac{114.6 \text{ Q}}{T} W(u)$$
 (2)

The value of the integral is given by the following series:

$$W(u) = -0.577216 - \log_e u + u$$

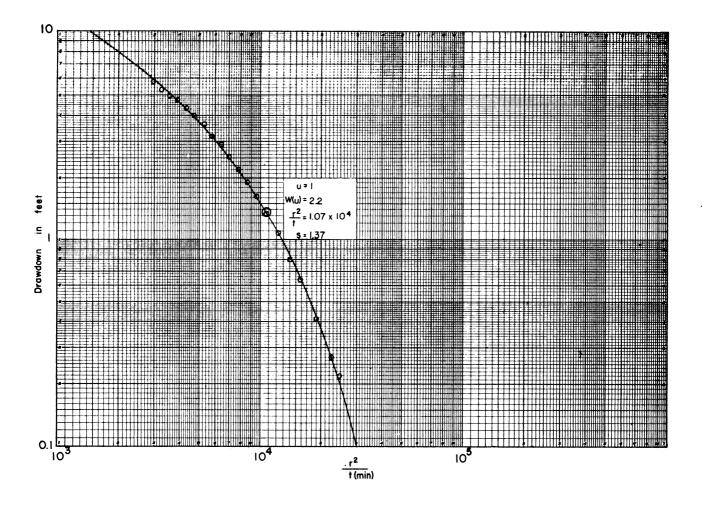
$$-\frac{u^2}{2.2!} + \frac{u^3}{3.3!} - \frac{u^4}{4.4!}$$
(3)
where  $u = \frac{1.87r^2S}{Tt}$ 

Values of W(u) for values of u between 9.9 and  $1.0 \times 10^{-15}$  are given by Wenzel (1942). The "well function of u" is plotted against u on loglog paper to form a type curve for determining the transmissibility and the storage coefficients of the formation tested. The observed draw-

downs are plotted versus  $\frac{\mathbf{r}^2}{t}$  on log-log paper

(Table 8). The graph of the observed data (Figure 6) is matched with the type curve by superposition, keeping the axes of the two

graphs parallel, and the values of  $\frac{r^2}{t}$ , s, W(u),



$$Q = 70 \text{ gpm}$$

$$T = \frac{114.6 \, Q \, W(u)}{8} = \frac{114.6 \, x \, 70 \, x \, 2.2}{1.37} = 1300 \, \text{gpd}_{/_{4+}}$$

$$S = \frac{uT}{1.87 r^2} = \frac{1 \times 1300}{1.87 \times 1.07 \times 10^4 \times 1440} = 0.000044$$

$$P = \frac{1300}{21} = 62 \text{ gpd/}_{ft^2}$$

Figure 6. Curve for Aquifer Test in the Tokio Formation with Well 175 Pumping 70 gpm and with Water-Level Measurements Made in Well 177.

and u are selected for any convenient point on the graph. The value of transmissibility is obtained from equation 2, and the value of the storage coefficient is then obtained from equation 4 (Figure 6). The coefficient of transmissibility near wells 175 and 177 is found to be 1.300 gallons per day per foot. The thickness of the water-bearing sand is 21 feet, therefore the coefficient of permeability is 62 gallons per day per square foot. The coefficient of storage was found to be 0.000044. This means that for each foot that the water level falls, 0.000044 cubic feet of water will be released from storage in each column the height of the aquifer with a base 1 foot square.

A fourth aquifer test was made on the city of Ashdown water supply wells in deposits of Quaternary age. Well 383 was pumped at 80 gallons per minute and water level measurements were made in an abandoned well 85 feet from well 383 (Table 9). The aquifer coefficients were determined by the Theis nonequilib-

TABLE 9

Data on Aquifer Test in Deposits of Quaternary Age with Well 383 Pumping 80 gpm with Water-Level Measurements Made in an Old Abandoned Well

--- O# # ---

Time since pumping started (minute)	r=85 feet= distance from pumping well  rs  t (min.)	Drawdown (feet) 8
2	3.6 x 10°	0.10
$\overline{4}$	1.8 x 10°	0.37
6	$1.2 \times 10^{\circ}$	0.62
8	$9.0  ext{ x } 10^{2}$	0.88
10	$7.2 \times 10^{\circ}$	1.12
12	$6.0 \times 10^{\circ}$	1.35
16	$4.5 \times 10^{\circ}$	1.75
20	$3.6 \times 10^{\circ}$	2.07
24	$3.0 \times 10^{\circ}$	2.32
30	$2.4 \times 10^{3}$	2.70
35	$2.1 \times 10^{\circ}$	2.90
40	$1.8 \times 10^{\circ}$	3.18
50	$1.4 \times 10^{-1}$	3.55
60	$1.2 \times 10^{\circ}$	3.85
70	$1.03 \times 10^{2}$	4.15
80	$9.0 \times 10$	4.36
90	8.0 x 10	4.60
100	$7.2 \times 10$	4.80
110	6.6 x 10	4.98
120	$6.0 \times 10$	5.10
130	5.55 x 10	5.20
140	5.15 x 10	5.30
150	4.8 x 10	5.40
200	3.6 x 10	5.95
300	$2.4 \times 10$	6.70
400	1.8 x 10	7.18

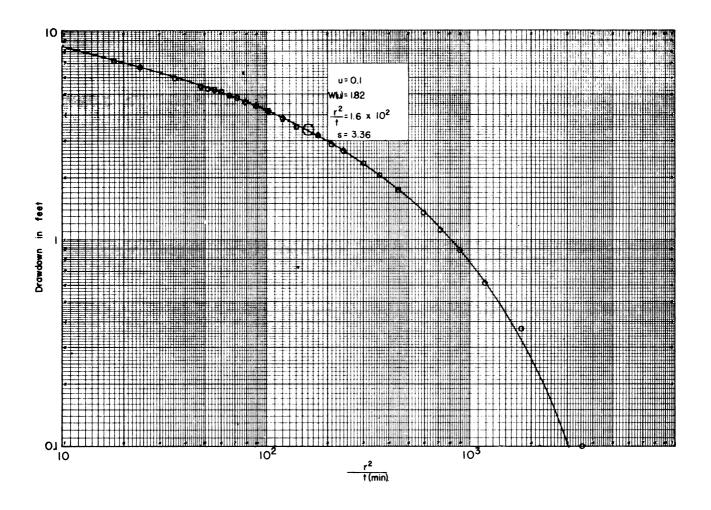
rium method. The coefficient of transmissibility was found to be about 5,000 gallons per day per foot. The aquifer thickness is about 30 feet, therefore, the coefficient of permeability is about 170 gallons per day per square foot (Figure 7). The coefficient of storage is found to be

0.0012. This storage coefficient is considered to be in the semi-artesian or leaky-aquifer range. It is too high for artesian aguifers and too low for water table aquifers. This is reasonable because the wells are screened with 20 feet of screen set in about 14 feet of sand at the bottom of the well with about 6 feet of the screen extending up into 11 feet of clay which overlies the sand. Another sand layer about 18 feet thick overlies the clay. The clay in these deposits is usually sandy and silty and water from the upper sand layer probably leaked through the clay into the sand layer tapped by the well. Under these conditions the coefficient of storage would increase with time and if this test could have been run much longer the storage coefficient would probably increase to about 0.1. Calculations of drawdowns made with the coefficient determined by this test would be greater than the actual drawdowns.

#### Recovery of Ground Water from Wells

When a well is pumped, the water level in the well, and for some distance from the well, declines and the water table or prezometric surface takes a form similar to that of an inverted cone, called the cone of depression. The greater the pumping rate of a well the greater the drawdown (difference between static and pumping water levels) and the deeper the cone of depression. The rate of discharge of a well per unit of drawdown is the specific capacity of the well. The Hope Light and Water Company well 309, in the Nacatoch sand, for example, yielded 280 gallons per minute with a drawdown of 118 feet, hence its specific capacity was 2.5 gallons per minute per foot of drawdown. The specific capacity of well 207 of the Hope Light and Water Company, in the Tokio formation, is 2 gallons per minute per foot of drawdown and for well 175 of the Ideal Cement Company, also in the Tokio formation, the specific capacity is 0.7 gallons per minute per foot of drawdown.

When a well is pumped, the water level drops rapidly at first and then more slowly, but it may continue to drop for several hours or even for many days or years. In testing the specific capacity of a well, it is important to pump the well until the water level declines only very slowly. When a pump is stopped, the water level in the well rises rapidly at first and then more slowly until it approaches or reaches its original position. Figure 8 is a graph showing the rate of recovery of the water level in well 207 of the Hope Light and Water Company after it had been pumped at a rate of 184 gallons per minute for 7.5 hours. After the water level had been



Q = 80 gpm  

$$m = 30 \text{ feet}$$

$$T = \frac{114.6 \text{ Q W(u)}}{6} = \frac{114.6 \times 80 \times 1.82}{3.36} = 5000 \text{ gpd/ft}$$

$$S = \frac{u}{1.87} \frac{T}{r/t} \frac{1}{(\text{days})} = \frac{0.1 \times 5000}{1.87 \times 1.6 \times 10^2 \times 1440} = 0.0012$$

$$P = \frac{5000}{30} = 170 \text{ gpd/ft}^2$$

Figure 7. Curve for Aquifer Test in Deposits of Quaternary Age with Well 383 Pumping 80 gpm with Water-Level Measurements Made in an Old Abandoned Well.

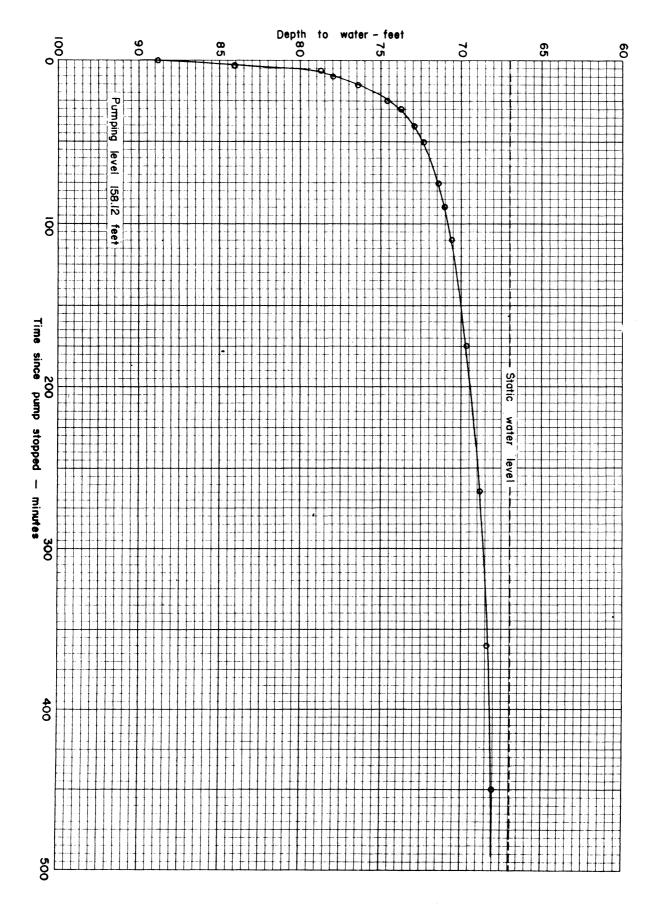


Figure 8. Recovery Curve for Well 207 of the Hope Light and Water Company.

allowed to recover for almost 8 hours it lacked 1.15 feet of returning to the original static level.

The character and thickness of the waterbearing materials have a definite bearing on the yield and drawdown of a well and, therefore, on the specific capacity of a well. Drawdown increases the height that the water must be lifted in pumping a well, thus increasing the cost of pumping. If the water-bearing material is coarse and uniform in size it will readily yield large quantities of water to a well with relatively small drawdown; if the water-bearing material is fine and poorly-sorted, as in most of the water-bearing formations in the area covered by this report, it will offer more resistance to the flow of water into the well thereby decreasing the yield and increasing the drawdown. Other things being equal, the drawdown of a well is inversely proportional to the transmissibility of the water-bearing material.

#### Utilization of Ground Water

There are many wells in the area covered by this report. All of the larger wells and approximately three-fourths of all the wells in the area were studied and the data obtained on these 397 wells are given in Table 11. The ground water withdrawn from these and other wells is estimated to average about 8.5 million gallons per day. The rates of use (in million gallons per day) for different purposes are: Public Supplies 3, Industrial 0.5, Domestic Use 3, Wastage 2, and Irrigation less than 0.1.

These figures are based on city pumpage records, industrial pumpage, not including industrial use from city supplies, per capita use of 50 gallons per day per person, including water for stock, of rural population, actual measurement of water from flowing wells going to waste, and acreage irrigated.

#### **Public Supplies**

There are 11 municipalities in the area that use ground water as a source of public supply. Eight of these supplies are described briefly in the following paragraphs.

Ashdown.—The city of Ashdown is supplied by 4 drilled wells, (383, 384, 385 and 386), in deposits of Quaternary age. The wells are 90 feet deep except well 385 which is 96 feet deep, and they are 10 inches in diameter except well 386 which is 8 inches in diameter. They contain approximately 20 feet of screen from about 70 feet to the bottom. The wells are equipped with electrically-driven turbine pumps with a reported capacity of about 80 gallons a minute

each. The output of well 383 was measured at 80 gallons per minute during the aquifer test at Ashdown.

The water is pumped from a 75,000 gallon ground reservoir with an electrically-driven centrifugal pump into a 50,000 gallon elevated tank. The water, which is distributed through the mains by gravity, is hard but suitable for most uses (see analysis well 386 in Table 12).

DeQueen.—The water supply for DeQueen is obtained from 3 wells (2, 4, and 5) that have a reported yield of 150 to 200 gallons per minute each. These wells yield water from sand and gravel in the Trinity group from a depth of 145 to 450 feet. The water is pumped from the wells with turbine pumps, driven by electric motors, into a 150,000 gallon ground reservoir. It is pumped from the ground reservoir into a 50,000 gallon elevated tank. The city has another well (3) that is not used except in emergencies because it pumps sand. It yields water from the same strata.

Analyses of the 4 wells (Table 12) indicate a total hardness of 58 to 180 parts per million.

Foreman.—The town of Foreman obtains its water supply from 2 dug wells (390 and 391) in deposits of Quaternary age. Their yields are 38 and 15 gallons per minute. The water is stored in an 11,000 gallon ground reservoir and a 50,000 gallon elevated tank.

An analysis for well 390 is given in Table 12.

Gurdon.—Gurdon is supplied with water from 2 wells (250 and 251) tapping the Nacatoch sand. They are 4 inches in diameter and are 245 feet deep. A third well (252) of the same depth is reported to yield water mixed with asphalt and is used only in emergencies. The wells are equipped with electrically-driven turbine pumps and the reported yield in gallons per minute are 60 for well 250, 50 for well 251, and 100 for well 252. The water is pumped into an elevated tank for distribution.

An analysis of water from well 250 is given in Table 12.

Hope.—Water is supplied to the city of Hope by 5 wells. Wells 208 and 209 are 1,480 and 1,500 feet deep, respectively, and yield water from the Tokio formation. Well 208, drilled in 1918, has 50 feet of perforated casing from 1,430 feet to 1,480 feet, and well 209, drilled in 1950, has 120 feet of screen from 1,380 feet to 1,500 feet. Wells 309, 310 and 314 are 620 feet deep and yield water from the Nacatoch sand. Well 314 is a gravel-wall well with 76 feet of screen and wells 309 and 310 have no screen. They draw

water through the open end of casing from about 230 feet of uncased hole. The city has 2 old Nacatoch wells (311 and 312) 620 feet deep that have no screen. They are not cased below 370 feet and have been abandoned. Wells 206 and 207, about 4 miles north of Hope, are on the former Army Air Force Proving Ground now owned by the city of Hope. These wells yield water from the Tokio formation. Well 206 is used to furnish water to a housing area near the well.

Water from the wells is pumped by electrically-driven turbine pumps into a ground reservoir and then pumped into an elevated tank for distribution through the city mains. The average rate of use is about one million gallons per day.

Analyses of water from wells 207, 208, 209, 309, 311 and 312 are given in Table 12.

Murfreesboro.—The city of Murfreesboro is supplied by one well (1) 8 inches in diameter yielding water from sand in the Trinity group. The depth of this well is not known. Water is stored in a 25,000 gallon ground reservoir and distributed from a 55,000 gallon elevated tank.

An analysis of water from well 1 is given in Table 12.

Prescott.—Prescott is supplied with water by 3 wells (152, 280, and 281). Well 152 is 1,070 feet deep, 8 inches in diameter, and yields water from the Tokio formation. Wells 280 and 281 are 250 and 325 feet deep and yield water from the Nacatoch sand. All of the wells are equipped with electrically driven turbine pumps. Storage capacity is 250,000 gallons.

Analyses of water from all 3 wells are given in Table 12.

Texarkana.—The city of Texarkana is supplied with surface water from Bringle Lake and ground water from 3 well fields. Two of these well fields are in Texas and are not described in this report. The Arkansas Station well field has 22 wells (397) ranging in depth from 40 to 50 feet and yielding water from deposits of Quaternary age. Their combined yield is about 700,000 gallons per day. They are equipped with vacuum pumps operated from a common power source.

A composite analysis (397) of water from all 22 wells is given in Table 12.

#### Industrial Supplies

Industries in the area use ground water at an estimated rate of 0.5 million gallons per day.

The largest industrial user is a cement company, located in southern Howard County, which uses about 0.2 million gallons per day. Numerous small industries such as lumber companies, ice plants, railroads, and a brick factory also use ground water.

#### **Domestic Supplies**

Many wells in the area furnish water for domestic use. The estimated rate for domestic use is 1.5 million gallons per day. The wells are dug or drilled and they range from 25 to 950 feet in depth. The drilled wells range from 2 to 4 inches in diameter. Water flows from some of the wells and some have sufficient pressure to pipe the water directly into the houses. Others have to be pumped and are generally equipped with small electrically powered pumps.

#### Irrigation Supplies

A small amount of ground water is used for irrigation in the area. Two wells located in the Red River valley, east of Texarkana, supply water for the irrigation of rice and the supplemental irrigation of dry crops. These wells each have yields of about 400 gallons per minute, and tap alluvium deposited by the Red River.

#### Water Wastage

In southwestern Arkansas there are two large areas and one small area in which water will flow from wells. The large areas are located in the Little Missouri bottoms in northern Hempstead County, southern Pike County, and northern Nevada County, and southwest of Nashville in southwestern Howard County. The small area is located in southern Sevier County in the Little River bottoms. In these areas there are numerous uncapped wells from which water is flowing and is not being used. The wells were drilled to furnish drinking water for people working in fields, for stock watering, and for use in connection with the lumber industry. Many of the wells are located a considerable distance from habitation. A total of 102 wells were found from which water was flowing and was being wasted. The rate of flow from these ranged from 1 gallon to 90 gallons per minute. averaging about 11. The total rate of flow from the observed wells was about 1.6 million gallons per day (see Table 10). Considering that all of the flowing wells were not found during the study it is likely that the total flow of water being lost was at a rate of about 2 million gallons per day.

TABLE 10
Water Wastage from Flowing Wells

County	Number of wells	Estimated flow (gallons per day)	
Hempstead	15	530,000	
Howard	32	360,000	
Nevada	10	60,000	
Pike		410,000	
Sevier		240,000	
Total	102	1,600,000	

It is estimated that in the Little Missouri bottoms the water level has been lowered as much as 8 feet during the last 20 years. On higher ground nearby, wells, from which water originally flowed, have ceased to flow and at present have to be pumped or have been abandoned.

Many of the flowing wells could be capped thus preventing the loss of water and the lowering of artesian pressure. The water from some of the wells might be used for supplemental irrigation.

#### Quality of Water

The chemical character of the ground water utilized in the report area is shown by 328 analyses of water listed in Table 12. The analyses were made by the U. S. Geological Survey, using standard procedures that are based on methods found in authoritative publications (Collings 1928 and Am. Public Health Assoc. 1946).

The analyses relate to the dissolved mineral constituents which determine the fitness of the water for municipal, industrial, agricultural, and domestic uses without reference to the sanitary quality of the sample. In most cases a single sample from a well is regarded as being representative of the chemical quality of the water because the concentration of the dissolved minerals of water in individual wells seldom changes appreciably. Exceptions to this generality include very shallow wells, where the concentration is modified by rainfall; wells tapping aquifers that are subject to salt-water encroachment; or wells recharged by streams that fluctuate in mineral concentrations.

#### Chemical Constituents in Relation to Use

The following discussion of the chemical constituents of ground water in relation to use was adapted from publications of the U. S. Geological Survey and is limited to only those constituents given in Table 12.

Iron.—Iron is dissolved from many rocks and soils and frequently from the pipes through which the water flows. The quantity of this constituent in ground water may differ greatly

from one locality to another, even though the waters are derived from the same formation. Waters with a low pH value (high in acidity) tend to dissolve iron readily. Large amounts of iron may interfere with efficient operation of silicate water softeners, and may prove more detrimental than either excessive hardness or dissolved solids. Reddish-brown stains on white porcelain or enameled ware and fixtures, and on clothing or other fabrics washed in water, are often caused by iron. The U.S. Public Health Service Drinking Water Standards, for interstate carriers, recommend a maximum of 0.3 part per million (ppm) of iron and manganese together in potable water. Many industries cannot tolerate any iron in their supply of process water; however, iron is readily removed by aeration and filtration. Unless otherwise noted, values of iron reported in Table 12 are for total iron. The iron content of water sampled in this area ranged from 0.01 part to 54 ppm. Of a total of 328, 241 samples contained less than 1.0 ppm; 59 contained 1.0 part to 5.9 ppm; and 16 contained more than 7.0 ppm iron.

Carbonate and bicarbonate.—Bicarbonate in natural water results from the action that dissolved carbon dioxide has on carbonate rocks. The carbonate radial (CO<sub>3</sub>) is generally not found in appreciable quantities in natural waters. Aside from the effect on the palatability of the water where present in excessive amounts, bicarbonate is of little significance in public water supplies.

Carbonate ranged from none in most samples to 98 ppm. Bicarbonate ranged from 3 parts to 773 ppm; most samples had less than 300 ppm.

Sulfate.—Sulfate, when present in sufficient quantities, combines with calcium to form hard, boiler scale, calcium sulfate. The sulfate ranged from 1.0 part to 560 ppm with most samples having less than 100 ppm. All but 11 samples had less than 250 ppm of sulfate, the U. S. Public Health Service recommended limit.

Chloride.—Since the chloride of calcium, magnesium, and sodium is readily soluble, chloride is normally present in most water. Sodium chloride (common table salt) in higher concentrations gives the water a salty taste, and also causes foaming and priming in boilers. Magnesium chloride in waters is believed to form hydrochloric acid by hydrolysis when heated, and such waters would be corrosive. The limit recommended by the U. S. Public Health Service is 250 ppm. The samples ranged from 3.0 parts to 3,850 ppm with all but 49 samples being below the recommended limit.

Nitrate.—Nitrate in water may result largely from the complete oxidation of nitrogenous materials. The quantities usually found in Arkansas waters are not sufficient to be objectionable. The nitrate content of water sampled from this area ranged from 0 parts to 560 ppm with 19 samples having more than 5.0 ppm and 6 samples having more than 40 ppm. Five of the 6 wells with a nitrate content greater than 40 ppm were shallow wells, ranging in depth from 22 to 60 feet; the depth of the other well is not reported. The high concentrations of nitrates are probably due to pollution by surface water, which has been in contact with fertilizers or other nitrate salts, leaking into the wells. Recent studies indicate that nitrates in water may be the cause of methemoglobinemia ("blue babies"). In a report published by the National Research Council, (Maxcy 1950) concludes that a nitrate content in excess of 44 ppm as NO<sub>3</sub> should be regarded as unsafe for infant feeding.

Hardness.—When choosing a water supply for either municipal, industrial, or domestic use, the hardness of the water is probably one of the most important factors considered. It is easily recognized by the increased quantity of soap required to produce lather, by the formation of the insoluble curd that is objectionable in all washing process, and by the deposits of insoluble salts formed when the water is heated or evaporated. In addition to its soap-consuming capacity, hard water is objectionable because of the formation of scale in boilers, water heaters, radiators, and pipes, with a resulting loss in heat transfers, possible boiler failure and loss in flow.

Calcium and magnesium are usually the chief causes of water being hard. Iron and aluminum likewise cause hardness, but are usually present in much smaller quantities than calcium and magnesium. When present as bicarbonate or normal carbonates, the hardness is called carbonate hardness; when present in the form of sulfates, chlorides, nitrates, and other soluble salts of calcium and magnesium, it is called non-carbonate hardness.

Water that has less than 60 ppm of hardness is usually rated as soft and suitable for most uses without further softening. Water with hardness ranging between 60 and 120 ppm may be considered moderately hard, but this amount does not seriously interfere with the use of water for many purposes, except in high-pressure steam boilers and in some industrial processes. Water with hardness ranging from 120 to 200 ppm is considered hard. Where the hardness is in the upper range, laundries and some indus-

tries may profitably soften their supplies. Water with hardness greater than 200 ppm needs some softening before it can be used satisfactorily for most purposes, though much harder water is used without treatment, especially for domestic purposes. Analyses of water from the wells sampled for this report show a range in hardness from 3 parts to 790 ppm, but generally less than 100 ppm, except for the aquifers of Quaternary age where it is usually more than 300 ppm.

Specific conductance.—The specific conductance of a water, a measure of its capacity to conduct a current of electricity, is dependent on the concentration and degree of ionization of the different minerals present. Although the conductance value does not give an indication of the relative quantities of the different constituents, it does give an indication of the total quantity of the dissolved minerals. The conductance is of particular use in those instances where the dissolved solids are not determined, since, for Arkansas waters, the dissolved solids value, expressed in ppm, is roughly half the conductance value, expressed in micromhos.

Hydrogen-ion concentration (pH).—The degree of acidity or alkalinity of water, as indicated by the hydrogen-ion concentration, or pH, is related to the corrosive properties of water and must be known for proper treatment for coagulation at water-treatment plants. The pH is the logarithm of the reciprocal of the hydrogen-ion concentration. For practical purposes, the pH scale ranges between the numbers 0 and 14. denoting various degrees of acidity or alkalinity. Neutral water has a pH of 7. Values below 7 and approaching 0 are increasingly acid, while values from 7 to 14 are increasingly alkaline. As the pH becomes less than 7.0, and especially less than about 5.0, the water becomes more corrosive to certain metal surfaces. The pH of the water sampled for this report ranged from 4.7 to 9.2, but in most cases the pH was above 8.0.

## GEOLOGIC FORMATIONS AND THEIR WATER-BEARING PROPERTIES

Mississippian and Pennsylvanian Systems

The Stanley shale, Jackfork sandstone, and Atoka formations crop out on the northern edge of the area covered by this report (Plate 2). These formations consist of folded and faulted hard shale and sandstone. The rocks contain little ground water but sufficient amounts usually can be obtained for domestic use. Veatch (1906) states that water-bearing sandstones

occur throughout this area but that it is quite possible to drill 1,000 feet or more without obtaining a good supply.

#### Jurassic System

Oil geologists have reported rocks of Jurassic age in deep oil tests in the southern part of the area covered by this report. They consist of dark shale, shaley limestone, sandstone, red beds, and anhydrite. These rocks have been overlapped by Cretaceous, Tertiary and younger deposits. They do not crop out and do not yield fresh water to wells. The section shown on Plate 1 does not go deep enough to show the Jurassic deposits.

#### Cretaceous System—

#### Lower Cretaceous (Comanche) Series

The Lower Cretaceous of southwestern Arkansas consists of the Trinity group and Kiamichi and Goodland formations. The Trinity is the only unit of early Cretaceous age that yields fresh water to wells in Arkansas. The Kiamichi formation and the Goodland limestone occur in a very small area near Cerrogardo in western Sevier County near the Arkansas-Oklahoma state line.

#### Trinity Group Undifferentiated

Character, distribution, and thickness.—The outcrop of the Trinity is shown on Plate 3. It includes the lowermost rocks of Cretaceous age in Arkansas. In the northern part of the area it overlies the Mississippian and Pennsylvanian rocks but in the southern part of the area it lies directly on truncated beds of Jurassic age. At its outcrop and for a short distance downdip the Trinity deposits indicate a shallow water, and occasionally a marginal, environment of deposition. The most common materials are clay, sand, gravel, and limestone. Some carbonaceous material is present in the clay and limestone beds.

The Trinity is divided into 6 units for this report. The lowermost unit is gravel. It ranges from 20 to 50 feet thick near its outcrop to a maximum known thickness of about 100 feet downdip. At its zone of contact with the Paleozoic rocks, the gravel is sometimes cemented with iron oxide. At other places it is composed of an unconsolidated mass of rounded pebbles and cobbles ranging in size from one-half inch or less to ten inches in diameter, intermixed and interbedded with sand and clay. The lower gravel unit represents the northern edge of the Lower Cretaceous outcrop and is seen as a very irregular band or as outliners capping many of the hills in the area. Apparently the gravels are

thinner to the west and overlapped to the east by younger beds.

The second unit of the Trinity is a limestone that overlies the lower gravel. It thins to the west and extends from the Cossatot River eastward to southeastern Pike County where it is overlapped by younger rocks. It ranges in thickness from 10 feet to about 40 feet and is composed of fossiliferous limestone and some green clay.

The third unit of the Trinity is mostly clay mixed with fine sand. It is thin at its outcrop and reaches a known thickness of nearly 400 feet downdip.

The fourth unit is a gravel lithologically similar to the lower gravel but finer grained. This gravel attains its maximum thickness of about 40 feet near the state line.

The fifth unit is a limestone and immediately overlies the upper gravel. It extends into Oklahoma on the west and to a point about four or five miles south of Murfreesboro on the east where it is overlapped by younger rocks. This upper limestone unit is about 60 feet thick and is composed of interbedded gray fossiliferous limestone and clay with small amounts of gypsum. Downdip the limestone lenses out and is replaced by anhydrite and red shale.

The uppermost unit of the Trinity consists of interbedded sand and red and gray clay and shale. Locally there are thin gravel lenses in this upper unit, but on the whole the sand is usually fine-grained and well sorted. The sand unit wedges out in western Hempstead County.

Near the southern part of the southwest Arkansas area, the Trinity group reaches a probable thickness of more than 2,500 feet.

Water supply.—The Trinity yields water to a part of the wells in southern Howard County and all of the wells in central Sevier County. Records of 37 wells are given in Table 11 and the locations of these wells are shown on Plate 3. The chemical character of water from 33 wells is given in Table 12.

The upper sand unit of the Trinity group is the principal source of water in the group. Flowing wells from this sand may be obtained at the lower altitudes in Howard and Sevier Counties. The largest yield observed from the upper Trinity is from well 34 in the Saline River bottoms which flows about 100 gallons per minute.

The quality of water from the upper sand is variable and similar to water from the overlying Woodbine formation. The Woodbine overlies the truncated edges of these sands and may be interconnected to some extent. The water in the Trinity would require some treatment for most industrial processes.

At and around DeQueen, the upper and lower gravels yield water of fair quality to wells. Wells in these gravels at DeQueen have reported yields of as much as 200 gallons per minute.

There are many dug wells in the outcrop area of the Trinity group that yield small amounts of water for domestic use.

#### Goodland Limestone

Character, distribution, and thickness.—The only outcrop of the Goodland limestone in Arkansas is at the extreme western edge of the area in the northwest corner of Little River County at the Oklahoma State line. The formation is probably less than 50 feet thick, being made up of hard sandy, light gray, fossiliferous limestone. The Goodland limestone crops out over a large area in Oklahoma.

Water supply.—The Goodland limestone does not yield water to wells in Arkansas.

#### Kiamichi Formation

Character, distribution, and thickness.—The Kiamichi formation overlies the Goodland limestone. It is probably less than 20 feet thick and is composed of alternating hard and soft, gray, fossiliferous marl and lenses of limestone. The outcrop of the formation also is limited to the extreme northwest corner of Little River County.

Water supply.—The Kiamichi formation does not yield water to wells in Arkansas.

#### Upper Cretaceous (Gulf) Series

The Upper Cretaceous includes all of the strata above the base of the Woodbine formation and below the base of the Midway formation of Tertiary age. The sequence of formations is shown on Table 5. Deposits of Gulf series rest unconformably on the sand of the Trinity group in the western part of the area but east of the vicinity of the Little Missouri River and in the vicinity of Arkadelphia they rest directly on rocks of Paleozoic age. The formations of late Cretaceous age crop out in parallel bands striking in general northeast-southwest directions. They crop out in a nearly continuous belt from the Ouachita River on the east to the Little River valley on the west. In the west and southwest part of the area, from the Little River valley to the Oklahoma State line. the Gulf sediments are covered by Quaternary terrace and alluvial deposits except for small

outliers near Foreman. In general the Gulf formations thicken from east to west.

#### Woodbine Formation

Character, distribution, and thickness.—The Woodbine is the oldest formation of the Upper Cretaceous series. It overlies unconformably the lower Cretaceous series in all of the area south of the outcrop area of the Trinity group and west of the Little Missouri River in Pike, Howard, and Sevier Counties. The thickness of the Woodbine formation ranges from a feather edge in eastern Hempstead County to 250 feet in Little River County.

The Woodbine formation is composed largely of reworked volcanic material with some sand, gravel, and some red and gray clay. The basal Woodbine consists of a lenticular gravel member which thickens eastward from a few feet to about 60 feet. The component materials range in size from pea gravel to large cobbles. Rounded and sub-rounded novaculite, chert and quartzite are most common with some pieces of well weathered igneous material.

The upper part of the Woodbine is made up chiefly of cross-bedded, yellowish sand containing reworked volcanic materials and some red and gray clay.

Water supply.—There are only a few wells in southwest Arkansas that yield water from the Woodbine formation. All of these are in southcentral Howard County. One well in southern Sevier County was reported to have been drilled into the basal gravels in the Woodbine formation. At this location the gravels were cemented and did not yield water and the well was abandoned. A well drilled for the former White Cliffs School at White Cliffs reached the gravel in the Woodbine. The water was reported to have been too salty to drink and the well was abandoned. No known wells tap the Woodbine formation in Hempstead County, probably because water is available from the overlying Tokio formation. It appears that the Woodbine might yield small amounts of water to wells in the northwestern part of Hempstead County.

Records of 10 wells are given in Table 11 and the locations of these wells are shown on Plate 3. Chemical analyses of water from 8 wells is given in Table 12.

The quality of water from the Woodbine formation is variable but usually high in sulfate and bicarbonate and low in chloride. The amount of chloride probably increases downdip as indicated by the reported "salt water" obtained in

a well at White Cliffs School. Hardness is generally low.

#### **Tokio** Formation

Character, distribution, and thickness.—The Tokio formation unconformably overlies the Woodbine formation west of the Little Missouri River and the Trinity formation and Paleozoic rocks east of the Little Missouri River. Tokio formation crops out in a northeastward trending band that is covered in stream valleys, and across most of Little River County by overlying Quaternary deposits. The outcrop area of the Tokio formation is shown on Plate 4. The outcrop area in general increases in width westward with an increase in formation thickness. The Tokio formation increases in thickness from about 50 feet near the edge of its outcrop to about 300 feet in Little River County, and probably continues to thicken westward and southward. Eastward it thins rapidly and is overlapped by the Brownstown marl a few miles east of the Little Missouri River in Clark County. The Tokio underlies all of southwest Arkansas south of the area in which it crops out and west of the Little Missouri River.

The Tokio formation is composed chiefly of cross-bedded sand and clay interbedded and intertongued with scattered carbonaceous material and some gravels. The basal gravel is the most persistent part of the Tokio formation but other lenticular beds of gravels occur higher in the formation. Sub-rounded pebbles of lightgray to brown novaculite are the most common constituents and range from pea size to six inches in diameter. The gravel is thickest in Hempstead and Howard Counties and thins eastward and westward. The dip of the gravel bed is not precisely known but is probably about 70 feet per mile south and southeast. The gravels lense out westward in Sevier and Little River Counties and are replaced by medium to fine sand.

The remainder of the formation consists largely of medium, cross-bedded, brown and gray sand with intertonguing and interbedded gray leaf-bearing, fossiliferous, clay, with varying amounts of lignite. The sand beds are usually poorly sorted. In southeastern Sevier County, southern Howard County, and western Hempstead County, there are three distinct sand beds within the Tokio as illustrated in Plate 6. They have not been named and will be referred to in this report as upper, middle, and lower sands of the Tokio. The upper sand occurs at the top of the formation. The middle sand occurs about 60 to 80 feet stratigraphically below

the upper sand. The lower sand, which is equivalent to the basal gravel in the outcrop area of the Tokio formation, occurs about 100 feet below the middle sand. To the east the clay layers, separating the sand beds, becomes thinner and in Hempstead County the three sand beds merge to form a massive sand which is prevalent over most of Hempstead County, southern Pike County, and northern Nevada County.

Much of the Tokio formation at the outcrop is sandy and weathers into a sandy soil. A mantle such as this is sufficiently permeable to allow percolation of surface and rain water into the sand.

Water supply.—The Tokio formation yields water to wells in southern Pike County, northern Nevada County, northern and central Hempstead County, southern Howard County, and southeastern Sevier County. Records of 165 wells are given in Table 11 and their locations are shown on Plate 4. Plate 4 also gives the altitude of the top of the Tokio by means of contours. From these contours, the approximate depth to the top of the Tokio may be determined by adding to the altitude of the land surface the values read from the contours where the top of the Tokio is below sea level or by subtracting from the altitude of the land surface, the values read from the contours where to top of the Tokio is above sea level. The upper water-bearing sands are immediately below the top of the formation.

Yields of up to 300 gallons per minute may be obtained in central Hempstead County. Flowing wells may be obtained in the bottom-land areas adjacent to streams. Wells flowing as much as 90 gallons per minute occur in these areas. Most of the water wastage from flowing wells given under utilization of water is from the Tokio formation.

Chemical analyses of water from 143 wells in the Tokio formation are given in Table 12. Isochlors for water from the Tokio are given on Plate 4. They are contoured on equal parts per million (ppm) of chloride. The approximate ppm chloride content for water from the Tokio can be ascertained in advance of drilling from the isochlors shown on Plate 4.

The quality of water from the sand and/or sands of the Tokio is fairly uniform, but there is wide variation in sulfate and iron content. In general the amount of chloride increases gradually downdip in the formation. However, the chloride increases abruptly in extreme western Clark and northwestern Nevada Counties.

The same abrupt rise in chlorides occurs in southern Sevier County (see Plate 4).

The iron content of the Tokio water is high near the center of northern Hempstead County and in parts of southern Howard County, ranging up to as much as 54 ppm.

Bicarbonate and sulfate are high indicating a moderately high sodium bicarbonate and sodium sulfate water. The water is generally soft with only a few exceptions.

#### Brownstown Marl (Restricted)

Character, distribution, and thickness.—The Brownstown marl unconformably overlies the Tokio formation and in Clark County overlies the rocks of Paleozoic age. The Brownstown marl crops out in an interrupted band which parallels and lies southeast of the outcrop of the Tokio formation. The chief exposures of the formation extend from the valley of the Little River northeastward to the Ouachita River valley. It is covered by Quaternary deposits in Little River County and in the stream valleys. The Brownstown marl is about 150 feet thick in Howard County, thinning slowly westward and more rapidly eastward, and it underlies all of the Cretaceous area south of its outcrop.

West of the Clark-Pike, and Clark-Nevada County lines the Brownstown marl is uniformly a gray marl with small quantities of sand and numerous large oyster shells. On weathering it turns a light tan. In Clark County the marl grades into a shallow water facies containing numerous sandy beds and sandy clay. In places the sand contains small quantities of chert pebbles which occur within hard calcareous sandstone lenses.

Water supply.—The Brownstown marl, being relatively impermeable, is not important as an aquifer in the Cretaceous area; however, it will yield small amounts of water for domestic purposes, mainly in Clark County where the amount of sand in the formation increases. The single analysis of water from the Brownstown in Clark County shows a highly mineralized water (see Table 12). Downdip from the outcrop the formation becomes less sandy and less permeable, and therefore the possibility of getting highly mineralized water is greater with increasing depth.

#### Ozan Formation

Character, distribution, and thickness.—The Ozan formation is unconformable with the underlying Brownstown marl. The basal unit of the Ozan, which crops out in southeastern Sevier County, southern Howard County, and western

Hempstead County, is known as the Buckrange sand lentil, a thin fossiliferous, very glauconitic sandy marl containing chert pebbles and phosphatic material. The Buckrange is not known to crop out in the eastern part of the Cretaceous area. The remainder of the Ozan formation, near its western outcrops in Arkansas, is composed of micaceous, fossiliferous gray sandy marl with a few beds of sand and sandy limestone. In Little River County, where it is largely overlain by deposits of Quaternary age, the Ozan changes facies from a sandy clay and marl to chalk and marl.

The Ozan formation in Clark County is similar in appearance and lithology to the Overlying Brownstown marl; however, the Ozan may be differentiated either by a glauconitic sand bed at the base, if present, or by the prevalence of smooth, convex oyster shells.

The Ozan is about 200 feet thick in Little River County, thinning to 150 feet in Howard and Hempstead Counties, and probably less than 75 feet in Clark County.

Water supply.—Records of 13 wells in the Ozan formation are given in Table 11 and the locations of these wells are given on Plate 3. Chemical analyses of water from 12 of the wells are given in Table 12. The Ozan formation yields small amounts of very highly mineralized water to wells in Clark County. This water is used only for domestic purposes and some of it is not suitable for that. Chlorides range from 32 parts per million to 2,100 parts per million. There is one well in southern Sevier County that yields water from the Ozan, but it contains 1,100 parts per million chloride and is not used. One well in Hempstead County yields water from the Ozan. It is least mineralized of the samples taken from wells in the Ozan formation containing only about 400 or 500 ppm total solids.

#### Annona Chalk

Character, distribution, and thickness.—The Annona chalk crops out near the southern edge of Howard County, the western edge of Hempstead County, and the northeastern and southwestern corners of Little River County. The Annona reaches a maximum thickness of about 100 feet at the outcrop in Little River County. It consists of massive white to gray, fossiliferous chalk, which in some places is sandy or marly. It is absent on the surface in the eastern half of the Cretaceous area but it may be present in the subsurface.

Water supply.—The Annona chalk does not yield water to wells in Arkansas.

# Marlbrook Marl (Restricted)

Character, distribution, and thickness.—The Marlbrook marl is exposed as a discontinuous band extending from the vicinity of Arkadelphia southeastward to the Saline River with one exposure in western Little River County. East of the Howard-Hempstead County line the Marlbrook lies directly on the Ozan and the contact is sharp; but to the west it overlies the Annona chalk and the contact is often transitional. The Marlbrook reaches a thickness of about 200 feet to the west and thins gradually eastward.

The Marlbrook marl is a uniform massive to thick-bedded, fossiliferous, chalky marl. It is blue gray in color when wet and when dry and exposed to weathering, it is a chalky white. Near the eastern edge of its outcrop area, the basal portion immediately overlying the Ozan formation is sandy and glauconitic.

Water supply.—The Marlbrook marl does not yield water to wells in Arkansas.

## Saratoga Chalk

Character, distribution, and thickness.—The Saratoga chalk unconformably overlies the Marlbrook marl throughout the Cretaceous area south of its outcrop. The unconformity marks a break in both lithology and fossil content. Thin outcropping bands of the Saratoga chalk extend southwestward from two miles west of Arkadelphia to the Little River bottoms. The formation is covered by Quaternary deposits in the stream valleys and in Little River County. The Saratoga chalk reaches a maximum thickness of about 60 feet.

The lower contact of the formation is characterized by a very thin glauconitic and phosphatic zone, with numerous fossil fragments. The greater part of the formation is white, very fossiliferous, sandy chalk.

Water supply.—The Saratoga chalk does not yield water to wells in Arkansas.

#### Nacatoch Sand

Character, distribution, and thickness.—The Nacatoch sand crops out as a wide, low ridge extending from the valley of the Little River in western Hempstead County northeastward to Arkadelphia. It is overlain by Quaternary deposits in the Little Missouri River valley, and Little River County. The outcrop area is shown on Plate 6. The outcrop averages from five to seven miles wide. The beds dip southeastward at about 40 feet per mile. The Nacatoch de-

creases in thickness from about 400 feet at the Little River to about 150 feet near Arkadelphia.

The Nacatoch sand may be divided into three lithologic units. The lower unit consists of interbedded gray clay, sandy clay, and marl; dark clavey fine grained sand; and hard irregular concretionary beds, and contains lenses of calcareous, fossiliferous, slightly glauconitic sand. The middle unit consists of a dark-greenish sand, which contains coarse grains of glauconite and weathers to lighter shades of green. It is generally fossiliferous where it is glauconitic. This unit also contains irregular concretionary beds. The upper unit is composed of unconsolidated gray, fine-grained quartz sand. Crossbedding is common in this sand. Locally the sand is massive and contains a few hard lenses and beds of fossiliferous sandy limestone. This upper unit is the principal water-bearing part of the Nacatoch.

Water supply.—Records of 150 wells in the Nacatoch sand are given in Table 11 and the locations of these are shown on Plate 5. Plate 5 also gives the altitude of the top of the Nacatoch by means of contours. From these contours, the approximate depth to the top of the Nacatoch may be determined by adding, to the altitude of the land surface, the values read from the contours where the top of the Nacatoch is below sea level or by subtracting, from the altitude of the land surface the values read from the contours. where the top of the Nacatoch is above sea level. The main water-bearing portion of the Nacatoch is immediately below the top of the formation. The Nacatoch sand yields water to wells in southwestern Clark County, north central Nevada County, central Hempstead County, northern Miller County, and extreme southeastern Little River County (see Plate 5). Flowing wells with yields of one or two gallons per minute, may be obtained in the lowest stream valleys in Clark and Nevada Counties. The Nacatoch sand will produce yields of 150 to 300 gallons per minute in Hempstead County and western Nevada County. In Miller County, eastern Nevada County, and Clark County yields are generally smaller and the water may contain considerable chloride.

Chemical analyses of water from 114 wells are given in Table 12. Isochlors for water from the Nacatoch sand are shown on Plate 5. They are contoured on equal parts per million (ppm) of chloride. The approximate ppm chloride content for water from the Nacatoch can be determined from the isochlors on Plate 5.

The Nacatoch sand yields a moderately soft alkaline water whose total mineral content differs considerably from place to place. Downdip 2 to 20 miles from the outcrop area, the water generally is too salty for most uses. The change between fresh and salt water is often very sharp going from less than 100 to over 1,000 ppm chloride in about 4 miles. In northern Miller County and in southern Clark and northeastern Nevada Counties the salt water occurs close to the exposed area of the Nacatoch. Veatch (1906) gives the following explanation for the occurrence of salt water near the outcrop area of the Nacatoch:

"Apparently the soluble salts have been dissolved and carried down by the continued inflow of rain water, and so the beds have been leached near the outcrop and the salts concentrated at a greater or less distance from it. The line between fresh and salt water is often very sharp, . . . where, from the same horizon fresh water is obtained in one well and salt water in another not 200 feet away."

A more likely explanation might be, since the Nacatoch is a marine deposit, that the sea water that was entrapped in the formation while it was being deposited has never been flushed out. There is not enough discharge downdip from the outcrop to allow very much movement of water through the formation, and the flushing process has been and still is very slow.

Sulfate and nitrate in solution are usually low, exceeding a combined total of 100 parts per million in only 3 of the 114 samples. Bicarbonate concentrations are moderate, varying from 2 to about 400 parts per million with the usual range of from 200 to 300 parts per million. Iron is low with very few exceptions.

#### Arkadelphia Marl

Character, distribution, and thickness.—The Arkadelphia marl overlies the Nacatoch sand south of its outcrop throughout the Cretaceous area. It is made up principally of blue to brownish clay-marl with occasional thin stringers of limestone. It is moderately fossiliferous. The formation's thickness varies from 100 to 150 feet.

Water supply.—The Arkadelphia marl does not yield water to wells in Arkansas.

Tertiary System—Paleocene Series Midway Formation

Character, distribution, and thickness.—The Midway formation overlies the Arkadelphia marl and is separated from it by an unconform-

ity which represents a considerable length of geologic time. The lower 30 to 50 feet consists normally of bedded calcareous clay containing at some places beds of dense gray limestone. The upper part of the formation consists of uniform gray clay with siderite concretions. The thickness ranges from about 20 feet at the outcrop to about 400 feet at the southern edge of the area. The Midway underlies all of the area south of the Arkadelphia marl outcrop.

Water supply.—The Midway formation does not yield water to wells in Arkansas.

Tertiary System—Eocene Series

#### Wilcox Formation

Character, distribution, and thickness.—The Wilcox formation overlies the Midway in northern Miller County, southern Hempstead County, southern Nevada County, and southern Clark County. It ranges in thickness from a few feet on its outcrop to about 600 feet in the southern edge of the area. The Wilcox is composed of a series of dark finely laminated beds of sand and clay containing much vegetable matter, mostly in the form of lignite.

Water supply.—The Wilcox supplies water to numerous shallow dug wells along the southeastern edge of the area except in Miller County where it is overlain by Quaternary deposits. There it is necessary to drill wells into the Wilcox. The yield of wells is small but large enough for domestic supplies. Many of the wells in southern Hempstead County, southern Nevada County, and southern Clark County are so shallow that they bottom well above the bottoms of the drainage channels and therefore tend to go dry during extended dry periods. Well locations are shown on Plate 3.

The hardness of 4 samples of water from the Wilcox formation ranged from 32 to 110 ppm which is considerably softer than the water in deposits of Quaternary age. The iron content was high in three of the samples but all were low in chloride. One sample had 86 ppm nitrate, which may indicate pollution from the surface.

# Quaternary System—Pleistocene and Recent Series

Alluvial and Terrace Deposits

Character, distribution, and thickness.—Alluvial deposits underlie the flood plains of the Ouachita, Little Missouri, Saline, Little River, and Red Rivers and their tributaries. They consist of sand, silt, gravel and clay, and range in thickness from 0 to about 50 feet. Terrace deposits underlie most of Little River County,

northeastern Hempstead County, and northeastern Nevada County. They consist of sand, clay and gravel, and range in thickness from a few feet in Nevada County to about 90 feet in Little River County.

Water supply.—Records of 17 wells that yield water from deposits of Quaternary age are given in Table 11 and locations of the wells are shown on Plate 3. Water is obtained from the alluvial and terrace deposits mainly in Little River County and northeastern Miller County. Yields of as much as 150 gallons per minute are reported at Ashdown, in Little River County, and yields of an estimated 400 gallons per minute are reported in eastern Miller County where the water is used for irrigation. Numerous shallow. dug, or driven wells yield water from Quaternary deposits in the stream valleys but the yields are small and the water (see Table 12) is frequently hard. Soft water is obtained by the city of Texarkana from local deposits of Quaternary age in the eastern edge of the city. The yield of 22 wells about 45 feet deep is about 700,000 gallons per day.

# SUMMARY OF WATER SUPPLY BY COUNTIES

#### Clark County

In northern Clark County ground water is available in small quantities from the sandstone and shale of Mississippian and Pennsylvanian age. Wells may obtain supplies at depths of 100 to 300 feet but it is quite possible to drill 1,000 feet or more without obtaining a good supply (Veatch 1906).

In central Clark County, the amount of sand increases in the Brownstown marl and it will yield small amounts of highly mineralized water to domestic wells (see analysis 213, Table 12). Downdip from the outcrop the formation becomes less sandy and the possibility of getting highly mineralized water is greater.

The Ozan formation, overlying the Brownstown marl, yields small amounts of very highly mineralized water to wells (see Table 12). The water is used only for domestic purposes and some of it is not fit for that. Chlorides range from 32 to 2,100 parts per million.

The Nacatoch sand yields water to wells in central and southwestern Clark County. Wells range in depth from less than 100 feet in the central part of the county, near the outcrop, to 568 feet in the southwestern part of the county. Reported yields of up to 100 gallons per minute have been obtained by the city of Gurdon and the Gurdon Lumber Company. Flowing wells,

with yields of one or two gallons per minute, may be obtained in the lowest stream valleys.

The Nacatoch sand yields a moderately soft alkaline water whose total mineral content differs considerably from place to place. In Clark County the Nacatoch water becomes salty close to its outcrop. The change between fresh and salt water is often very sharp going from less than 100 to 1,000 parts per million chloride in about 4 miles.

The Tokio formation is present along the western edge of Clark County but it contains water high in chloride content. Well 53 in Clark County and well 52 in northeastern Nevada County have 1,200 and 920 parts per million chloride respectively. Any wells drilled to the Tokio in western Clark County will likely yield salt water unless they are located very close to its outcrop area.

The record of wells and chemical analyses of water for Clark County are given in Tables 11 and 12. The locations of wells are given on Plates 3, 4, and 5.

## **Hempstead County**

In northern and central Hempstead County water is obtained from sands in the Tokio formation. Wells vary in depth from 16 feet in the northwestern part of the county in or near the Tokio outcrop area to 1,500 feet at Hope. Reported yields of up to 300 gallons per minute are obtained at Hope. Flowing wells from the Tokio may be obtained in northeastern Hempstead County in the Little Missouri bottoms. Well 72 (Table 4) had a measured flow of 90 gallons per minute.

In central Hempstead County the Nacatoch sand yields water of good quality to wells. Yields of 150 to 300 gallons per minute have been reported at Hope. Depth of wells range from a few feet in the outcrop area (Plate 5) to about 700 feet in the southwestern part of the county.

In the extreme southern part of Hempstead County water is obtained for domestic use from shallow wells in sands in the Wilcox formation.

The record of wells and chemical analyses of water for Hempstead County are given in Tables 11 and 12. The location of wells are given on Plates 3, 4, and 5.

# **Howard County**

In northern Howard County ground water may generally be obtained in small quantities from sandstone and shale. Sands in the upper part of the Trinity group supply highly mineralized water to wells in central Howard County south and southwest of Nashville. Flowing wells may be obtained at the lower altitudes that flow from 1 to about 50 gallons per minute. The sands of the upper Trinity may be interconnected to the basal gravel of the overlying Woodbine formation. Flowing wells with similar yields and quality of water are obtained from the Woodbine in southern Howard County.

There are many wells in southern Howard County that derive their water from the middle and lower sands of the Tokio formation. In the extreme southern end of the county wells yield water from the upper sand of the Tokio. Wells vary in depth from a few feet in the outcrop area (Plate 4) to about 500 feet in the southern part of the county.

The quality of Tokio water is generally good. Exceptions are high iron and sulfate content in local areas. Chloride increases gradually downdip in the formation (Plate 4).

Tables 11 and 12 give the records of wells and chemical analyses of water from wells in Howard County. The locations of wells are given on Plates 3 and 4.

# Little River County

Little River County is almost completely covered by deposits of Quaternary age. These deposits yield water to wells from sand and gravel about 30 to 90 feet below the surface. The cities of Ashdown and Foreman obtain water from wells in these deposits. The water is generally hard.

The only Cretaceous deposits that yield fresh water are the Nacatoch sands in southern Little River County. Wells range in depth from about 100 to 365 feet.

A Trinity well (37) has a flow of highly mineralized water (see Plate 3). This well was drilled for an oil test. It is not known at what depth the flow of water is from. Any well drilled into the Trinity group in Little River County will likely yield salt water.

The Woodbine formation is reported to yield salt water at White Cliffs. The Woodbine overlies the Trinity except for a small area in northwestern Little River County. The Woodbine would likely yield salt water in Little River County.

The Tokio formation is reported to have been tested at various places in Little River County but it yields only salt water except for one well (204) in the western part of the county. It is a shallow dug well, 22 feet deep, where the Quaternary covering the Tokio is thin. An analysis of water (204, Table 12) from this well shows 102 parts per million chloride, hardness of 92 parts per million and iron of only 0.7 parts per million. Nitrate is 130 parts per million, which probably indicates that the water is contaminated by surface water.

Tables 11 and 12 give the records of wells and chemical analyses of water from wells in Little River County. The locations of wells are shown on Plates 3, 4, and 5.

## Miller County

In northern Miller County the Nacatoch sand is the only Cretaceous aquifer that yields fresh water to wells. The chloride content of water from the Nacatoch increases from about 100 parts per million in the extreme northern part of the county to over 3,000 parts per million near Texarkana (Table 12). Wells in the Nacatoch vary in depth from 250 feet in the northern part of the county to about 1,000 feet at Texarkana.

Sands of the Trinity group, Woodbine formation, and Tokio formation would likely yield salt water in Miller County as they do further to the north in Little River County.

The Wilcox formation supplies water to domestic wells near Texarkana where water is not available from overlying deposits of Quaternary age and where the Quaternary is missing (see Plate 2). Wilcox wells have small yields but generally the quality of water is better than that from the Quaternary (see Table 12).

Deposits of Quaternary age cover the northern part of Miller County northeast of Texarkana. Sand and gravel in these deposits can generally be reached at depths from 40 to 80 feet. These deposits in the Red River valley east of Texarkana yield about 400 gallons per minute of poor quality water to irrigation wells (see Table 12).

Tables 11 and 12 give the records of wells and chemical analyses of water from wells in Miller County. The locations of wells are shown on Plates 3 and 5.

#### **Nevada County**

In extreme northern Nevada County, in the bottoms of the Little Missouri River, flowing wells yield water from the Tokio formation. These wells vary from about 225 feet to 600 feet deep. The city of Prescott and the Arkansas Louisiana Gas Company each has a well yield-

ing water from the Tokio. These wells are 1,070 feet and 1,217 feet deep. They yield water that has a high chloride content (see wells 152 and 205, Table 12). Any wells drilled into the Tokio more than 10 miles south or east of Prescott would likely yield salty water (see Plate 4).

The Nacatoch sand crops out across the northern end of Nevada County. It yields water of good quality to dug wells in and near its outcrop and to drilled wells at and south of Prescott. Nacatoch wells will flow a few gallons per minute in the lowest stream valleys. Five flowing wells in T. 11 S., R. 20 W., yield water with more than 500 ppm chloride (see Plate 6). Wells range in depth from a few feet in the outcrop area to over 600 feet about 15 miles south of Prescott.

In the southern part of the county dug and drilled wells yield water from the Wilcox formation. The town of Rosston has the only known municipal well yielding water from the Wilcox in the area.

The record of wells and chemical analyses of water in Nevada County are given in Tables 11 and 12. Locations of wells are shown on Plates 3, 4, and 5.

## **Pike County**

In the part of Pike County north of the Trinity outcrop (Plate 3), domestic supplies can usually be obtained above 300 feet. In the region of the Trinity outcrop (Plate 3) very little ground water is available. The Trinity group is here a calcareous clay of practically no water bearing value (Veatch 1906).

South of the Tokio formation outcrop (Plate 4) flowing wells can be obtained along the Little Missouri River valley and some of its tributaries. The depth of wells, in the Tokio here, range from a few feet in the outcrop to about 250 feet near the Little Missouri River. The quality of water from the Tokio is good.

The records of wells and chemical analyses of water from Pike County are given in Tables 11 and 12. The locations of wells are shown on Plates 3 and 4.

## **Sevier County**

In northern Sevier County north of the Trinity outcrop area, water may usually be obtained from Mississippian and Pennsylvanian rocks in sufficient quantities for domestic use.

In the vicinity of DeQueen, the two gravel units of the Trinity group, yield water of fair quality to wells at depths of from 145 feet to 450 feet. Trinity wells at other places throughout central Sevier County vary in depth from a few feet to 415 feet.

The Tokio formation yields water to wells in southern Sevier County. The wells range in depth from about 100 feet, near the outcrop area, to about 400 feet near the southern tip of the county. Many Tokio wells flow a few gallons per minute in the Cossatot River bottoms.

The records of wells and chemical analyses of water from Sevier County are given in Tables 11 and 12. The locations of wells are shown on Plates 3 and 4.

#### WELL LOGS

Listed on the following pages are the drillers' logs of 10 wells in southwestern Arkansas. The descriptions of the materials given in the logs are as they were reported by the driller.

The geological classification of the logs was made by the senior author. Many of the logs are either too generalized or too incomplete to permit an accurate geological classification, so that the formation boundaries are not exact and their thicknesses may not be accurate.

# WELL 4

Owner: City of DeQueen

Location: SW14 NE14, sec. 30, T. 8 S., R. 31 W.

Altitude: 450 ft. (approx.)

Driller: --

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age	V	
Surface material	12	12
Gravel	6	18
Trinity group		
Limestone	4	22
Clay, red	8	30
Limestone		37
Clay, blue	9	46
Limestone		49
Clay, blue		52
Limestone		67
Clay, yellow		77
Limestone		78
Clay, yellow	3	81
Limestone		90
Gravel and clay, red		108
Limestone		113
Slate, blue	32	145
Sand and gravel		160
Limestone		167
Gravel and clay	12	179
Limestone		180
Gravel and clay	12	192
Sandstone	1	193
Gravel	58	251

# **WELL 205**

# WELL 206-Continued

Location: SW 1/4 NE 1/4, sec. 3, T. 12 S., R. 23	
Docation. 5 W 74 NE 74, Sec. 5, 1. 12 5., 16. 25	W.

Altitude: 320 ft. (approx.) Driller: H. Cox

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age		
Surface clay	40	40
Arkadelphia marl		
Shale	183	223
Nacatoch sand		
Sand rock	. 8	231
Sand		248
Sand and gravels		305
Sand and gravels (shell and rocks		325
Shale Shale	10	335
Broken sand rock		350
Sand		450
Saratoga chalk and Marlbrook marl	. 200	
Gummy shale	. 290	765
		765
	20	100
Ozan formation	0.5	050
Gummy shale		850
Shale and boulders	0.0	900
Gumbo	30	930
Grownstown marl		
Shale and boulders	. 30	960
Shale	40	1,000
Gummy shale	<b>29</b>	1,029
Sand	6	1,035
Shale	10	1,045
Sand	6	1,051
Gummy shale	. 19	1,070
Sand	. 15	1,085
Tokio formation		
Gumbo	35	1,120
Shale		1,140
Gumbo	18	1,158
Sand	14	1,172
Gumbo	6	1,178
Sand	. 39	1,217

# **WELL 206**

Owner: Hope Light and Water Co.

Location: SE¼ NE¼, sec. 3, T. 12 S., R. 23 W.

Altitude: 340 feet

Driller: Layne-Arkansas Co.

	Thickness (feet)	Depth (feet)
Arkadelphia marl		
Yellow clay		35
Blue shale	72	107
Nacatoch sand		
Fine sand	5	112
Rock	4	116
Rock, not so hard	6	122
Break		125
Fine sand		131
Rock	6	137
Fine sand	12	149
Hard rock	5	154
Fine hard sand	6	160
Hard spot		162
Fine sand	<u> </u>	164
Sand with hard spot	<b>2</b>	166
Sandy shale	<b>4</b>	170
Rock		171
Fine sand	20	191
Rock	1	192
Break	1	193

WELL 200-Continue	<del></del>	
·	Thickness (feet)	Depth (feet)
Rock	1	194
Fine sand		216
Finer sand		234
Rock		236
Breaks with rocks	_	242
Rock		245
Fine black sandRock		263 265
Rock Fine black sand		205 270
Rock	_	271
Sandy shale		275
Rock	. 1	276
Sandy shale with hard spots	14	290
Rock		291
Sandy shale		295
Rock		296
Sandy shale		307
Rock		308
Hard shale	42*	450
Saratoga chalk Tough clay or shale	113	563
Marlbrook marl		
Hard spot	. 8	571
Hard tough shale		763
Ozan formation		
Hard spot	. 6	769
Soft shale	. 14	783
Hard spot		786
Hard shale		813
Soft shale		826
Hard spot	-	829
Hard shale Shale and boulders	. 64	893 901
Tough shale		901
Boulders		903
Tough shale		904
Soft gummy shale		955
Brownstown marl		
Rock	. 1	956
Hard shale	. 95	1,051
Rock	. 1	1,052
Sandy shale		1,060
Rock		1,061
Soft shale		1,074
Hard spot		1,075
Soft shale Hard spot		1,085
Tokio formation	. 4	1,087
Soft shale	22	1,109
Shale and boulders		1,115
Sand	_	1,122
Sand (last 5 ft. not so good)		1,155
Sandy shale	. 16	1,171
Hard rock	. 1	1,172
Sand	15	1,187
Rock	2	1,189
Tough shale *Error either in thickness or depth of material	14	1,203
*Error either in thickness or depth of materia depth is correct; therefore the thickness should	ai. It is ass I read 142 an	umed that d 192 feet.

**WELL 207** 

Owner: Hope Light and Water Company Location: SE¼ SW¼, sec. 6, T. 12 S., R. 24 W.

Altitude: 340 feet

Driller: Layne-Arkansas Company

	Thickne (feet)	
TT 1 -1	10 14 16	10 24 40

# WELL 207—Continued

•	Thickness (feet)	Depth (feet)
Shale	46	86
Hard spot	. 2	88
Soft sandy shale Hard rock	. 7 4	98 98
Hard rock Soft sandy shale	. 3	102
Vacatoch sand	. 0	102
Rocks with soft spots	. 7	109
Shale	_	116
Rock	. 2	118
Shale	. 3	121 122
Rock Sandy shale		129
Rock		132
Soft sandy shale	. 8	140
Rock	. 2	142
Sandy shale	8 . 3	150 153
Hard rock Hard shale		161
Hard spot		163
Shale		166
Sandy shale		176
Hard spot		177
Soft sandy shale Hard spot		214 218
Fine black sand	15	230
Rock		234
Sandy shale	. 11	24
Rock		246
Blue shaleRock		256 257
Blue shale		276
Rock		277
Shale with hard spot		282
Hard spot	. 1	283
Shale Tough clay		296 322
Clay, some softer	10	332
Clay, some softer Tough shale	. 60	392
Boulders	. 4	396
Shale	_	414
Hard shale Shale and boulders		416 428
Gumbo	. 18	443
aratoga chalk and Marlbrook marl		
Hard shale	173	616
Shale (not so hard)	. 22	638
Hard shale	_	660
Hard gumbo		668 682
Sandy shale Soft shale	4.4	693
Fine sand		695
Soft shale	. 35	730
Tough shale		748
Sandy shale	15	760
ozan formation	10	776
Tough shale Sandy shale		77( 774
Shale		845
Hard shale		875
Sandy shale	. 7	882
Hard tough shale		887
Sandy shale Soft shale		900 925
Hard tough shale		932
Soft shale		944
rownstown marl		
Shale	. 16	960
Sandy shale		979
Boulders	. 1	980
Sandy shale	- 3	983
Lough Clay	. Z 11	985 996
		1,003
Sandy shale Tough clay Soft sandy shale Hard shale	. 11	99

	Thickness (feet)	Depth (feet)
Soft shale	1	1,004
Hard shale		1,009
Tough shale		1,043
Boulders	1	1,044
Sandy shale	23	1.067
Lignite	2	1,069
Sandy shale		1.077
Tough shale	17	1,094
Tokio formation		,
Water bearing sand	49	1,143

# **WELL 209**

Owner: Hope Light and Water Company Location: SE¼ SW¼, sec. 28, T. 12 S., R. 24 W.

Altitude: 353 feet

Driller: Holliday Well Company

•	Thickness (feet)	Depth (feet)
Surface soil	14	14
Midway formation		
Sand and clay	. 61	75
Shale and small rocks	156	231
Arkadelphia		
Shale and large boulders	139	370
Nacatoch sand		
Sand, shell, and large boulders all		
this section mixed conglomerate	234	604
Shale blue, gumbo	46	650
Sandy shale	. 52	702
Saratoga chalk and Marlbrook marl		
Shale and lime	. 102	804
Chalk	236	1,040
Ozan formation		,
Gray shale	. 17	1,057
Sand and sandy shale		1,100
Chalk and shale	. 160	1,260
Brownstown marl		,-
Sandy shale	40	1,300
Gray shale	. 80	1,380
Tokio formation		,-
Sandy shale	. 70	1,450
Water sand-fine grain	40	1,490
Black shale	10	1,500

# **WELL 299**

Owner: Arkansas-Louisiana Gas Company Location: SW4 NE4, sec. 3, T. 12 S., R. 23 W.

Altitude: — Driller: H. Cox

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age		
Surface clay	40	40
Arkadelphia marl		
Shale	80	120
Broken sand rock	. 5	125
Sticky shale		225
Nacatoch sand	100	
Sand and rock	3	228
Sandrock	-	234
Sand	-	245
Sandrock		248
Sand	-	256
Sandrock		262
Sand and gravel		272
		275
Sand and gravels (hard shells)		284
Sand and gravels (hard shell rock)	) 17	301

## **WELL 300**

Owner: Arkansas-Louisiana Gas Company Location: SW1/4 NE1/4, sec. 3, T. 12 S., R. 23 W.

Altitude: --Driller: H. Cox

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age	Annual Control of the	-
Surface clay	25	25
Arkadelphia marl		
Shale	35	60
Slicky shale	75	135
Gumbo	20	155
Shale		222
Nacatoch sand		
Sandrock	9	23
Water sand		248
Rock	•	250
Sand and gravel		260
Rock		262
Sand and gravel		272
Water sand		280
Rock		286
Sandrock		300
Rock		303
Sand		318
Rock		320
Sand and gravels		325

## **WELL 337**

Owner: Henry Mack

Location: NW 1/4 NE 1/4, sec. 23, T. 13 S., R. 26 W.

Altitude: 270 feet Driller: F. E. Wells

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age		and the second s
River dirt	64	64
Gravels	1	65
Arkadelphia marl		
Blue dirt	165	230
Nacatoch sand		
Rock	4	234
Sand		265
Rock—below 265		

## **WELL 308**

Owner: Hope Brick Company

Location: NW 1/4 SW 1/4, sec. 27, T. 12 S., R. 24 W.

Altitude: 350 feet.

Driller: Holliday Well Company

THE PARTY OF THE P	Thickness (feet)	Depth (feet)
Surface clay	30	30
Midway formation		
Shale	48	78
Boulders	0′ 2″	78′ 2″
Arkadelphia		
Shale	261′ 10″	340
Nacatoch sand		
Sand—fair	24	364
Rock	2′ 6″	366′ 6″
Sandy	1′ 6″	368
Rock	<b>2</b>	370
Sand		372
Rock	1	373
Sandy	2	375
Rock	1	376
Sandy	8	384
Rock	5	389
Sandy		410
Rock	4	414
Sandy	11	425
Rock		425′ 4″
Sandy	1′ 8″	427
Sandy shale	7	434
Rock	0′ 8″	434′ 8″
Shale and shell	5′ 4″	440
Rock	3	443
Sandy shale	22	465
Shale and boulders	15	480
Shale	117	597

## **WELL 386**

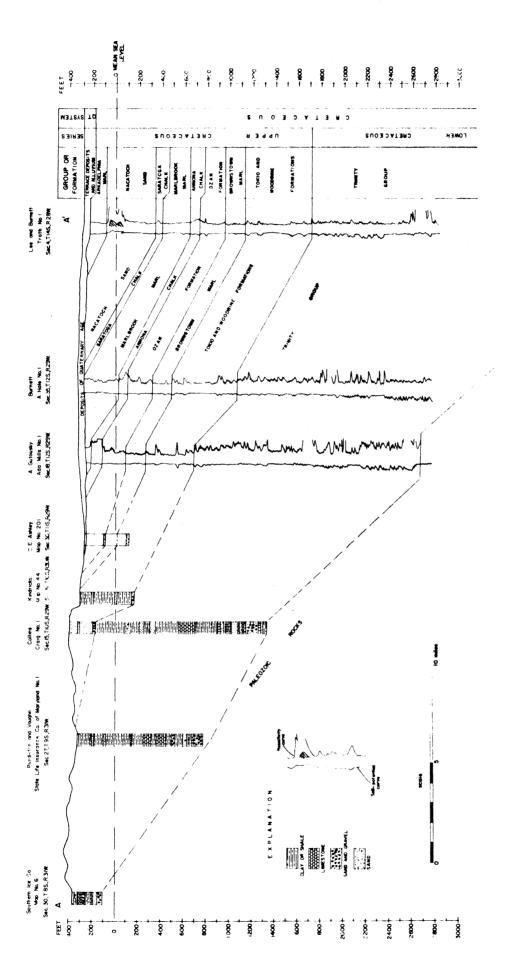
Owner: City of Ashdown Location: NE¼ SW¼, sec. 32, T. 12 S., R. 29 W. Altitude: 330 feet Driller: Shirly

	Thickness (feet)	Depth (feet)
Deposits of Quaternary age		-
Clay	8	8
Sand	30	38
"Joint clay"	9	47
Sand	18	65
Clay		76
Sand		90
Marlbrook marl		
Shale		below 90

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   a reconnaissance: U. S. Geol. Survey Circ. 241.
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t I. Generalized geologic section from DeGueen to southeastern Little River County

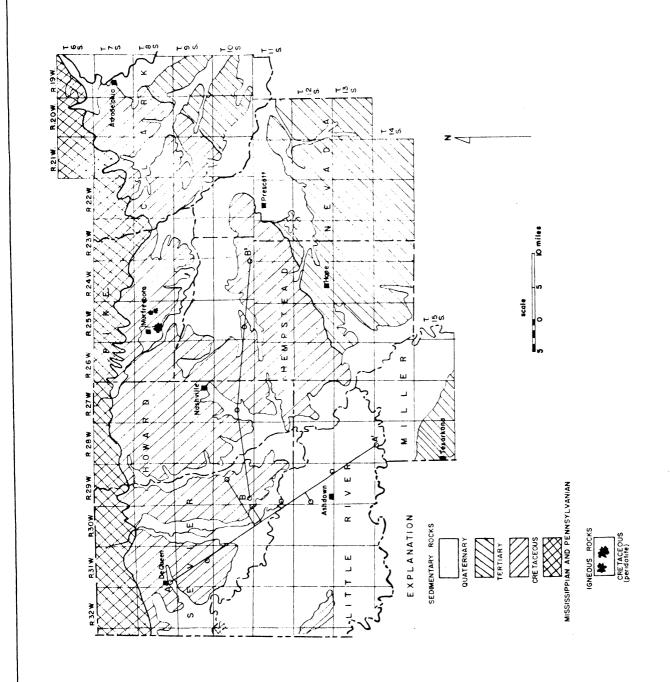
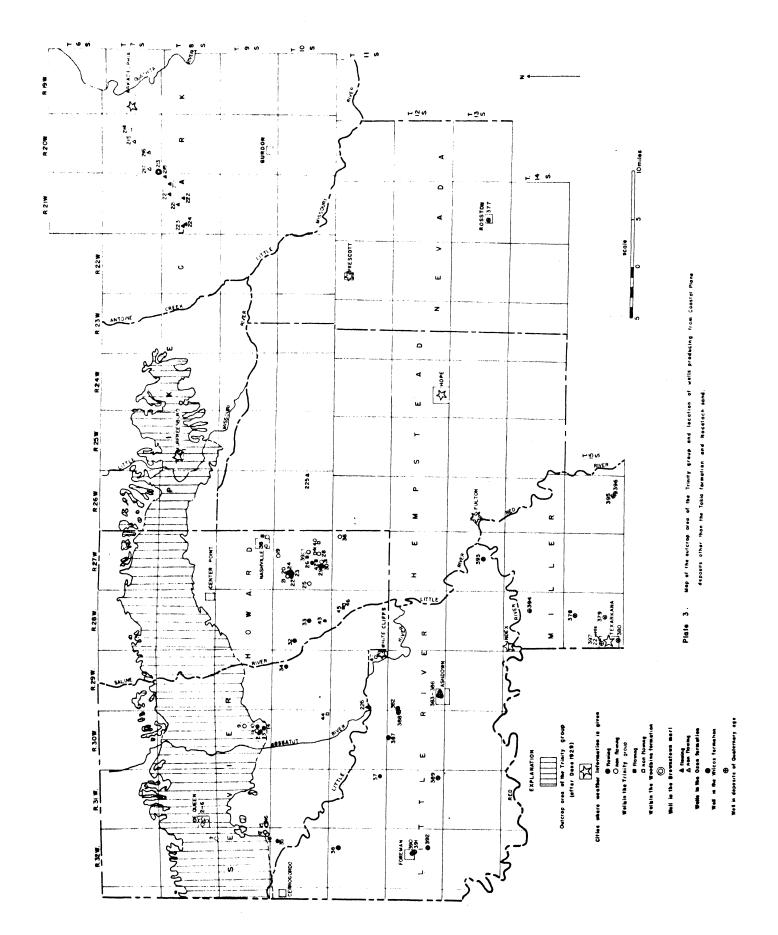
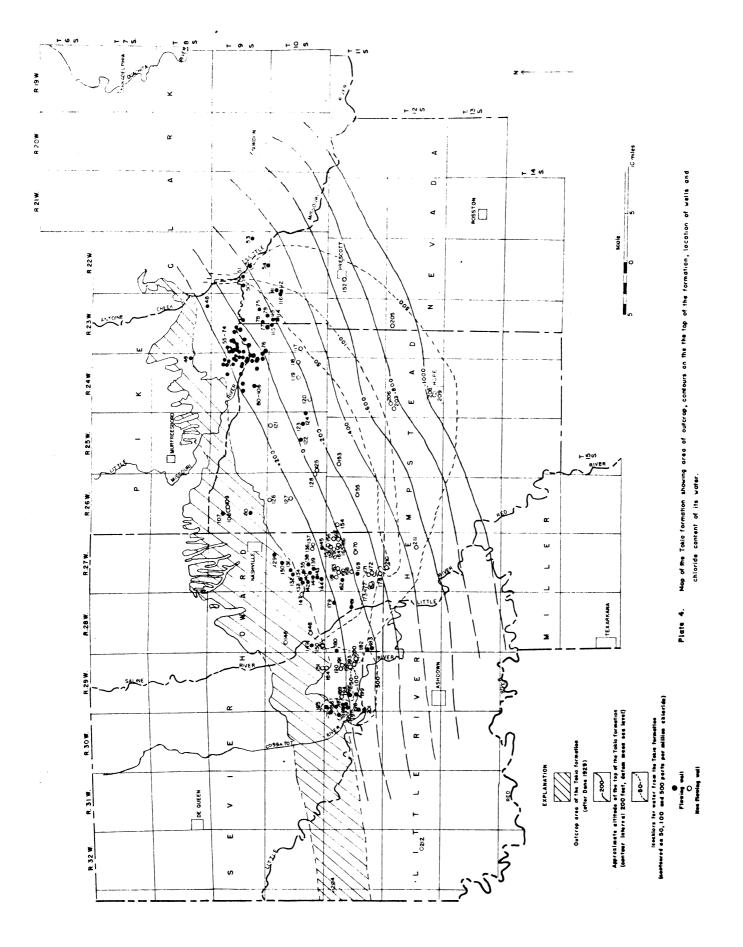
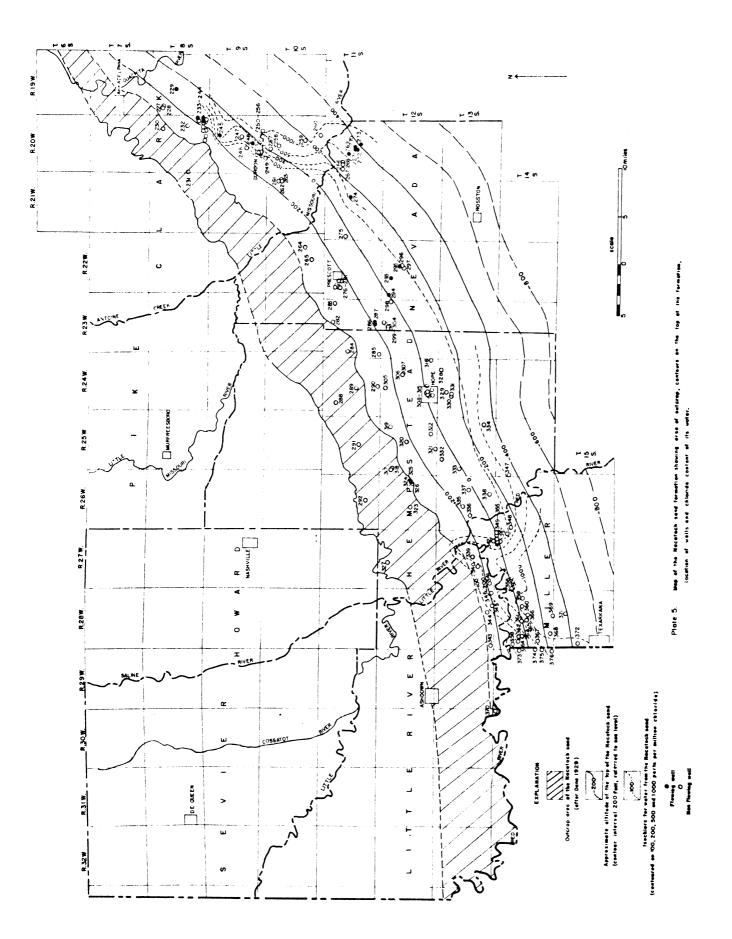
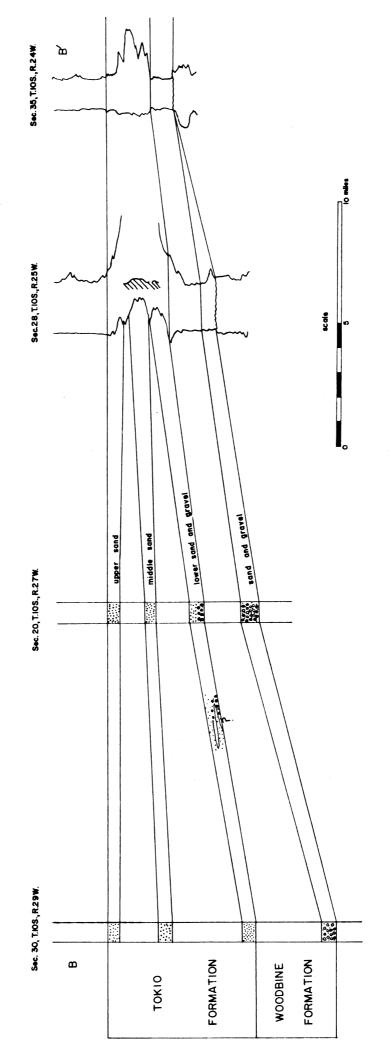


Plate 2. Geologic map of southwestern Ark, (after Dane 1929)









Diagrammatic cross section of sands of the Tokio and Woodbine from southeastem Sevier County to eastern Hempstead County Plate 6

TABLE 11. RECORD OF WELLS IN SOUTHWESTERN ARKANSAS

					!	-		nse.	
					Oil test well			Water reported too salty to use	
	Kemark			₹	nity nity_Oil			orted too	
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Water level feet below	- Contract		flowing almost dry flowing flowing flowing	flowing flowing flowing flowing flowing flowing flowing	nn se in se in se in se in se	ine	20 flowing flowing dry hole	84	
RECORD OF WELLS IN SOUTH WESTERN ARKANSAS Approx. Cushr Wate Date sittude Doyl diam. Get		20R	flow flow flow flow flow	flowing 1.5R 1.6R flowing flowing flowing flowing		flow		flow 10R	flowing flowin
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TABLE 11. RECORD OF WELLS IN SOUTHWESTERN ARKANSAS-Continued

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RECORD OF WELLS IN SOUTHWESTERN A	Driller et
TABLE 11.	ŧ

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Owner	W. T. Schooley	Ark. Agri. Exp. Sta. Ark. Agri. Exp. Sta. Hope Brick (G. Hope Lt. & Water Co. Hope Lt. & Water Co.	Hope Lt. & Water Co. Hope Lt. & Water Co.	Southern re Co. Hope Lt. & Water Co. Hope Basket Co. Francet McWilliams	Earnest we'r minams Mrs. P. M. Gilbert H. B. Gilbert	Hope Air Fleid R. E. Long Guernsey Public School	M. E. Fatrick Lee Cheetum E. R. Calhoon D. F. Goodlit	T. E. Edwards W. M. Dillard	Hope Flooring & Lbr. Co.	G. Anthony Lot. Co. N. B. Coleman Roy Woods	Spring Hill School Dist. Hope School District	40. Fac. K. K. Henry Mack Jesse Smith	Smith Place H. H. Horton, Jr.	J. B. Davis R. G. Reed J. B. Davis	Temple Brothers J. B. Davis	E. E. Faxton Mary Jefferson R. G. Reed	R. G. Reed R. G. Reed R. G. Reed	R. G. Reed R. G. Reed	R. G. Reed R. G. Reed A P Cox	R. G. Reed J. D. Harris	water Clement E. J. Miller Texarkana Const. Co.	T. L. Sinham M. A. Parker Mrs. Welloce Dunn	D. L. Thompson H. C. Shuffield	Mrs. Dan Gillham Chester Stanley Williams Canda Co	Kelly Budd Roy Steuber Wildred Pencentet	J. B. Richardson Walter Clement	Ed McKee H. T. Ham
County	Hempstead	Hempstead Hempstead Hempstead Hempstead	Hempstead Hempstead	Hempstead Hempstead	Hempstead Hempstead	Hempstead Hempstead	Hempstead Hempstead Hempstead	Hempstead Hempstead Hempstead	Hempstead Hempstead	Hempstead Hempstead		Hempstead Hempstead Hempstead		Little River Little River		ad		Miller Miller	Miller Miller	Miller	Miller Miller	Miller Miller	Miller	Miller Miller	Little River Miller Willer	Bowie, Texas Bowie, Texas	Bowie, Texas Bowie, Texas
Location number	12S24W-4ndd1	-14bac1 -14bac2 -27cbc1 -28cdc1	-28cdd2 -28cdd3	-33aac1 -33bcd1 -36aaa1	12S25W-7abb1 -7abd1	-15dbc1 -33ddc1 -34bc16	12S26W-21aac1 -24abc1 -24bda1	-24cal 12S27W-4ddel 13S24W-2adal	-4dcb1 -9bda1	13S25W-5abd1 -18aab1	-25ceb1 13S26W-16dda1	-23abb1 -35bc1	13S27W-22cab1 -28bbd1	-51acul -35ddd1 13S28W-31cdb1	-34adb1 -35cba1	14S25W-7bbc1 14S26W-7caa1	14S27W-18aa1 -18ad1 -18bb1	-1bbb1 -2aaa1	-2aab1 -2add1 -7cbc1	-12abi 14S28W-7ccai	-19001 -140bd1 -16dae1	-17bbc1 -17cba1 -18cc1	-10aac1 -21aaa1 -22bda1	-30aca1 -32deb1	14S30W-1daa1 15S28W-10aab1	East. Bowie Co., Tex.	East. Bowie Co., Tex. East. Bowie Co., Tex.
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TABLE 11. RECORD OF WELLS IN SOUTHWESTERN ARKANSAS-Continued

Remarks																					22 wells pumped with common pump in 1951
										See log											22 well pum
B		Z C	A A		۵	A)	S	<u></u>	S	S	Ω	Ω	A,	K	8	-	۵	Ω	Ë	II.	PS
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Date of measurement		145	1-28-51					7-24-51	7-24-51		1::		-51	94.	94		*********				7-25-51
Water level feet below land surface?		125R	35.30		1\$R	******		41.65	41.40			** * * * * * * * * * * * * * * * * * * *	10R	•	•		***************************************			:	26.95
Gasing flam. (fn.)		9			::	:::	10	10	10	œ	::	!		1	:		!	:	:	:	o O
<b>1</b>		2	270 150	A B E	8	35	S	8	96	90	S	65	22-60	!	:	23	30+	40±	<b>%</b>	20	40-50
Approx. akitude (feet)	a tion		340		1	:	330	99 99	330	330		:	342	1	:	\$	-	260	*****	:	300
Date frilled	form	1928		Quat		-	1912	1915	1944	1950		-		1	!	!			********		
Driller	Wilcox formation		E. Williamson W. K. Elliot	Deposits of Quaternary age					Howard Hipp	Shirly		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1					1				
Owner		Town of Rosston Roy Stanber	Miss Osteen Phillips W. K. Elliot		W. D. Cason	Ada Mills	Ashdown Water Works	Ashdown Water Works	Ashdown Water Works	Ashdown Water Works	S. M. Beck	Ada Mills	Lloyd Wright	Town of Foreman	Town of Foreman	Magnolia Pipe Line Co.	Smith Place	E. J. Miller	Luther Lowe	Shelly Frank	Texarkana Water and Sewage Co.
County		Nevada	Miller		Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Little River	Miller	Miller	Miller	Miller
Location number		13S21W-21cda1'	-27aba1 -32cab1	AND THE RESIDENCE OF THE PARTY	11S32W-2cc1	12S29W-6ccc1	-32cabi	-32cab2	-32cab3	-32cab4	12S30W-3bab1	-12aaa1	12S31W-36bcc1	12S32W-15dbb1	-15dbb2	-26bca2	13S27W-22cab-	14S28W-14cbd2	15S26W-26dcc1	-35aab1	15S28W-19ddd1
Map number		(377)	(372) (380)	-	(381)	(385)	383	384	382	(386)	(381)	388	(389)	(330)	391	(392)	(393)	(394)	(395)	(368)	(397)

1 Well number in parenthesis Indicates that analysis of water is given in table 12.

19 Water Pacific Proported.

10 Wei PS., public reported.

10 Wei PS., public analy 7. I industrial; D. domentic; 1/r., irrigation; N. not used.

10 Wei PS., public analy 7. I industrial; D. domentic; 1/r., irrigation; N. not used.

17 No means: T. 8 S., R. 25 W., sec. 17, 5W.y., SW.j., NB.y., See well aumbering system page 6.

TABLE 12. CHEMICAL ANALYSES OF WATER FROM WELLS IN SOUTHWESTERN, ARKANSAS

Car. Biener. Car. Biener

Depth Date of the bonate bonate Suifate Chierles Nitrate Non-Combination of Computation (feet) collection 'P' (Feet (Co.) (HCOs) (SOu) (C)) (G) (MCOs) Total Carbonate Microrbonal Pill

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Micrombon)		326	308	230	328	- C	21.7	GF.	7C	2.650	282	283	325	150	846	552	900 900 900 900 900 900 900 900 900 900	557	906	429	5.36	0 P	x t	2 2	825	495	512	324	213	1,050	04.6.7.			s Sic	(3X)	70	9.7	1 20	951			186	6.07°C	1.280	3,820	4,760	1.550	3:	150	86.4	9.75	9.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00	E	50e 80e	62.5	108	1 5 2 5 3 6	523	24.2 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0	50.0	269	6.69	15.	6.5.9	127	
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(foet)			145	251	291 450	076	250+		150±	22	140±	140±	150±	290	265	391	265	300	365	408	1006	914	454	415	510	340	415	:	100	00/			100	364	570	286	347	460	468			32	00-400	+00+	400 H	35	164	700∓	230	228	1005 1005	39	200	902	H 076	500+	225±	300	412	+ 90%	460	85	170	140	202	
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Location number		8S25W-17acc1	8S31W-30acd	-30acd	-30acd3	30dbc	9S27W-36bbb	9S30W-23bba	-26cbc	-27daa	-27dca	-27dcc	-35ppp	9S32W-36caa2	US27 W -3bdd1	-6acc1	-8cdb1	-8dac1	-8dcb1	-19dbb1	-21ddc	DD877-	-28cdb	-28cdc	-28dac	10S28W-18acc1	-21daal	0S29W-11bdc	10002 W -20001	11S31W-36bac1			9S27W-35bab	10S27W-22bdc1	-26adc	-2/acc	0S28W-33ade	1S28W-11bbc	-11bbd		THE STATE OF	9223 W-35000	9S22W-19aab1	-20abb	-21dec	Jedcz-	9S23W-7ccc1	-16bce	-16dd	-17bbc	-17bac	-18bdb	-18cac	-18ccc	-18dbe	-19beb1	-22baa	-26cdc1	-31cac1	-34bac	-34ddc1	9524 W - 10dda	-13dec	-1 4aac	-21 <b>88</b> C	
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Hardness a Total Ca	1	: -
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Sulfate (SOc.)	peq	
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r KOM Iron (Fe)	mati	88 24 4 4 4 4 4 4 6 6 6 6 6 6 6 6 6 6 6 6
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TABLE 12. CHEMICAL ANALYSES OF WATER FROM WELLS IN SOUTHWESTERN, ARKANSAS -- Continued

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Remarks		Upper sand		Middle sand	Transacti	o block same																																	The second secon														
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1 ABLE 12. CHEMI			Ideal Cement Company			Ideal Cement Company	ideal Cement Company O. A. Mitchell	Guy McCullough	Lewis Walk Edwin Polk	J. K. Polk J. W. Roberts	Milton Garrison	Loyd Garrison	Ben Savage	J. C. Walker	Mrs. w. C. Sutton C. I. Brewer	Vaughn	Dierks Lumber Company Rainh Penny	Frank Moore	C. E. Ashley Mrs. Phyna					Hope Light and Water Co.												Baswell Farms			Earl Manning N. W. Crawford			Grover Wells R H Herron			R. E. Allen		City of Gurdon F. E. Knight	Marvin Anderson	William Wells	Melvin Gates		Gurdon Lumber Company Gurdon Lumber Company	
Location number		11S27W-29ddb1	-30caa1	11S28W-2ddd1	-7bbb1 -14cad1	-30bcc1	-30cdb1 11S29W-2baa1	-6baa1	-7bbd1	-7dcc1	-11cabl	-11dcb1	-14ddc1	-17ade1	-1 /ubbi -18aba1	-18bdb1	-18cbb1	-24babl	-30bbb1	12daal	-12aaa1	12S23W-3acd3	12524 W-55001	-28cdd1	1252 (W -400c1 -23bda1	Marie to the country and the property of the property of	7S20W-31cbb1	V 2000000 0 11111000000		7S20W-22bbb1 -28cbc1	-30dcb1	8S21W-1acd1	-3ddd1	-9ddal	-10aber -18aaer	-18aac2 10S26W-24abc1	11S29W-30bbb2		8S19W-7aba1 8S20W-11abc1	-19ccel	-25cdd1	-25dbd1 -35caa1	-36bdd2	-21 dec1	-22abd1	-32dce1	-33bbc1 -33ddb1	-34abc1 -34dab1	-35caal	-9aaa1	-22cdd1 -35bcb1	10S21W-12bba1 -12bdd1	10S22W-23dcb1
Map		172	173	179	96 56	182	183	185	187	188 80 80 80 80 80 80 80 80 80 80 80 80 8	190	191 193	193	194	196	197	199	200	201 203	203	204	202	208	503	211	10 C 10 C 10 C 10 C 10 C 10 C 10 C 10 C	213			215	217	218	220	221 33	223	224 225	226		227 230	231	233	234 239	242	246	247 248	243	253 253	254 255	256	823	523 560 560	261 2 <b>6</b> 3	264

Iron (Fe)

Temperature

Date of collection

Depth (feet)

Location number

Remarks

Car- Blearbonate bonate Sulfate Chloride Nitrate Non- Conductance
(COs) (HCOs) (SOs) (Cl) (NOs) Total Carbonate (Micrombas) pH

	<b>4</b>	
	City No. 4	
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Ward   Walter Clement	Map	Location number	Owner	Perth (feet)	Pate of collection	Tompera-	lron (Fe)	Car- bonate by (COs)	Blent- bonnie S (HCO <sub>3</sub> ) (	Sulfate (SO <sub>1</sub> )	Chlorble (Cl)	Nitrato	Hardness as CaCOs Non- Total Carbonate	Sacos Same Con	On Specific Conductance to (Micrombos)	7	)temarka	
HSSPW.Ardehil Walfrer Chement					ž	cator	h sand	1-Con	tinue	~						:		
Horizon   Language	359	14S28W-14cbd1	Walter Clement	900	8-8-51		.10	1	516	:	200	ų.			.750	7.7		1
The control of the	36.	-14cbd1	E. J. Miller	\$ \$ +1 +	8-9-51	: 6	.17		92		410	دن (			95	oc oc		
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Carbon   H. C. Shuffield   G00+ R-S.51	365	-21aaa1	D. L. Thompson	450±	8-9-51		27		316		565	; c			500	ic o		
Mar. Dan Gillham   Stock   Webster Stander   Webster   Webster	366	-22bda1	H. C. Shuffield	₹009	8-8-51	!	.17		218		.670	2.0			100	- t -		
14530W-33dzb    Chester Stanley   Chester   Chester Stanley   Chester Stanley   Chester   Chester Stanley   Chester St	367	-30aca1	Mrs. Dan Gillham	200	8-8-51		.19		368		602	1.9			400	- 00		
Fast Bowie Co., Tex.   All three   Percent	308	-32dcb1	Chester Stanley	88	7-26-51	74	.12		245		,480	1.0			006	8.6		
Bast, Bowle Co., Tex.         Williard Pencantet         950         7-25-51         4         98         21         3850         15         11200         82           Bast, Bowle Co., Tex.         JB Redardson         700         8-51         10         8         6         24         6         24         0         2.720         8.15           Bast, Bowle Co., Tex.         JB Redardson         800         8-51         10         10         3         260         8         11         0         120         8.15         10         10         8.15         10         11         0         11         0         12         8         11         0         12         0         12         4         12         4         2         26         2         12         0         13         12         15         0         12         0         13         12         10         12         10         12         10         12         10         12         10         12         10         12         10         12         10         12         10         12         10         12         10         11         10         10         10         10         10 </td <td>971</td> <td>15C98W - 10cab1</td> <td>Nelly Budd</td> <td>3/6</td> <td>8-8-51</td> <td>28 28 28</td> <td>1.5</td> <td></td> <td>999</td> <td></td> <td>57</td> <td>1.7</td> <td></td> <td></td> <td>116</td> <td>7.4</td> <td></td> <td></td>	971	15C98W - 10cab1	Nelly Budd	3/6	8-8-51	28 28 28	1.5		999		57	1.7			116	7.4		
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SS2W-Zicda	375	East. Bowie Co., Tex.		200	8-8-51		.07		452		260	900			150	e ox		
18821W-21ctal   Town of Rosston   200   124-46   186   188   26   48   12   110   0   250   81   18528W-10aab2   Roy Stetch Philips   270   7-28-51   18   0   8   2   45   86   35   25   274   66   66   65   65   20   65   66   65   65   65   65   65   6	376	East. Bowie Co., Tex.	1	360	8-8-51	65	.10		306		290	4.1			330	28.2		
1882 W-2 tetal   Town of Rosston   124-46   106   0   148   26   45   12   110   0   274   66   158   128						=	0	a t	u o				and the same of th	İ			441	
1832W-2cct	377	13S21W-21cda1	Town of Rosston		1-24-46		190		148	26	8.4	1.2		0	006	- x		
Part   Part	00 C	15S28W-10aab2	Roy Steuber	30	7-26-51	!	1.8		<b>∞</b>	23	45	98		55.	27.4	99		
1532W-2cc1	380	-2/aba1 -32cab1	W. K. Elliot	270	7-28-51	1	0.0		=======================================	0.5	ر- ر تن ب	ಟ		0	203	6.6		
1532W-2ect   W. D. Carson   30   7-24-51   0.01   0   11   17   136   510   400   391   1540   5.5   12202W-6ect   Ashdown Water Works   35   7-26-51   0.02   0   11   17   167   3.9   455   1270   5.1   12502W-6ect   Ashdown Water Works   50   8-7-51   66   0.02   0   234   13   34   7   17   207   15   17   17   17   207   15   17   17   207   15   17   17   17   207   15   17   17   17   207   15   17   17   17   207   15   17   17   17   207   15   17   17   17   207   15   17   17   207   15   17   17   207   15   17   207   15   17   207   15   17   207   15   17   207   15   17   207   17   207   17   207   17   207   17   207   17   207   17   207   17   207   17   207   17   207   2	3	Topogo	A. W. Dillow	201	1-62-01		6.7		93	7	5.5	<b>x</b> :	1	0	152	1.1		
18.529W-Gecet   W. D. Garson   30 7-24-51   0.94   0 11   17 136 56.0 400 391 15.10 5.5     18.529W-Gecet   Ad Mils   18.5   1	The state of the s				ė	0	of Qua	artern	arya	e e								
12529 Weet   Academia   Academi	381	11S32W-2ccc1	W. D. Carson	30	7-24-51		600			17	136	560			540	5.5		
12830W-35ah   S.M. Beck   S.	386	12323 W -56661	Ada Mills	e e	7-26-51		50.5			r- (	167	3.9			270	8.1		
12831W-38bert	387	12S30W-3bab1	S. M. Beck	25	7-95-51	9	<u>۽</u> 1			2.5	34	7.7			711			
12832W-Jscheel   Lloyd Wright   126.60   125.51   102   134   12   124   212   351   241   102   1.0	388	-12aaa1	Ada Mills	36	7-26-51		7			-15	101	0.1			D 1 0 5	<del></del>		
12S32W-25dbal Town of Foreman	383	12S31W-36bcc1	Lloyd Wright	55-60	7-25-51		120			. 6	194	919.1		·	000	د د		
22 7-25bal Marnolia Pipe Line Co. 22 7-25-51 02 0 276 5 33 3.8 243 17 550 7.8 13827W-22bal Smith Pipe Line Co. 40 4 8-51 3.0 0 328 30 7.2 15 10 22 14 488 82 13 828 W-14cbd E.J. Miller E.	300	12S32W-15dbb1	Town of Foreman		7-11-46		.03			1 to	59	399			212	0.4		
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15528W-1960d1 Texarkana Water & Sewage Co. 40-50 172 188 38 30 7.2 9.6 349 80 711 8.1 8.1 8.1 8.2 8.2 8.3 9.6 349 80 711 8.1 8.1 8.2 8.2 8.3 9.6 849 87 80 654 0 1.840 7.9 8.3 8.3 9.0 9.6 849 8.3 8.3 9.0 9.0 849 87 80 87 80 8.3 8.3 9.3 9.3 9.3 9.3 9.3 9.3 9.3 9.3 9.3 9	303	13S27W-22cabl	Smith Place	30+		!	5.2			22	13	1.0			488	20.00		
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